

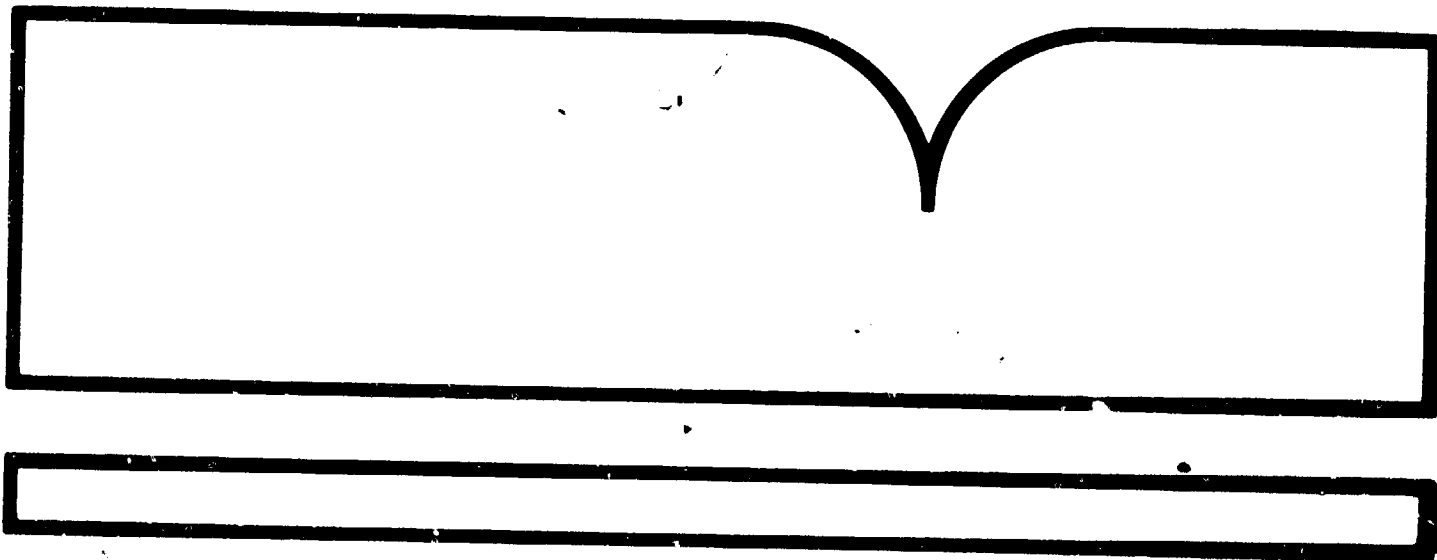
Hazard Index Calculation for May 31, 1984
Microburst at Erie, Colorado

(U.S.) National Oceanic and Atmospheric
Administration, Boulder, CO

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NOAA Technical Memorandum ERL WPL-155



HAZARD INDEX CALCULATION FOR MAY 31, 1984
MICROBURST AT ERIE, COLORADO

R. A. Kropfli

Wave Propagation Laboratory
Boulder, Colorado
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PREFACE

The NOAA Wave Propagation Laboratory in Boulder, Colorado, has been contracted by NASA Langley Research Center to compute hazard indices from dual-Doppler radar data taken during the PHOENIX II boundary layer experiment. This final report summarizes the results of these calculations.

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HAZARD INDEX CALCULATION FOR MAY 31, 1984 MICROBURST AT ERIE, COLORADO

ABSTRACT. The two x-band Doppler radars, operated by the NOAA Wave Propagation Laboratory, were used to collect high resolution data within a small, benign-looking microburst during the PHOENIX II boundary layer experiment. The lowest 2.5 km of the microburst was observed throughout its development and dissipation over a 15 minute period. These observations presented an excellent opportunity to compute a quantitative threat to a hypothetical aircraft whose flight track would carry it through the microburst. This hazard index is based on the kinetic energy loss to the aircraft that would be produced by the microburst; it is a function of the vertical air motion, horizontal spatial derivatives of the wind field, and the assumed aircraft air speed and direction. Indices were computed and plotted for all eight volume scans and peak values were observed to be sufficiently high to present a significant hazard to an aircraft even though the virga-produced microburst was visually unimpressive.

1. INTRODUCTION

During the PHOENIX II convective boundary layer experiment conducted during May and June of 1984, the two x-band Doppler radars operated by the NOAA/Wave Propagation Laboratory documented the evolution of a small microburst. This microburst was located 30 km NW of Denver's Stapleton International Airport and was one of a family of microbursts that were evident in the area on that day. Twenty minutes prior to this event another microburst was believed to have been the cause of significant damage to UAL Flight 633 upon takeoff from Stapleton (NTSB, 1985).

The dual-Doppler radar synthesis performed with this data set is of excellent resolution, both temporally and spatially, because of the short radar baseline (13 km) and scan volume repetition cycle (100 s). A description of the wind fields synthesized on a 150 m grid at 100 m intervals is given by Kropfli (1986) and is provided in Appendix I.

The Hazard Index, an aircraft-specific measure of the threat for a given wind shear condition (Bowles and Targ, 1988), has been computed for this data set. Contour plots of the Hazard Index have been plotted for the lowest horizontal plane for each of eight volume scans, and they are included here in Appendix II. Four flight directions were assumed. These data have also been recorded on two ASCII formatted tapes that are included along with this report.

2. DISCUSSION

The Hazard Index, as defined by Bowles and Targ (1988), is the loss in available excess thrust-to-weight ratio of an aircraft due to vertical air motions and horizontal windshear. It is expressed as follows:

$$F = (V/g)du/dx - w/V$$

where V is the airspeed, g is acceleration due to gravity, du/dx is the horizontal gradient of the horizontal wind along the direction of the flight path, and w is the vertical component of the wind. Their analysis shows that the potential rate of climb can be expressed as

$$dh/dt = \{[T - D]/W - F\}V$$

where T is the aircraft thrust, D is the drag, and W is the aircraft weight. From this expression the Hazard Index is seen to be the tangent of the difference between the actual glide slope from the desired glide slope caused by vertical air motion and horizontal wind shear. Values of F in the neighborhood of .10 to .15 are believed to be reasonable threshold levels above which an aircraft is in danger.

For this study four flight directions were assumed for the 31 May microburst; *toward* the north, east, south, and west and the corresponding Hazard Indices are expressed as FN , FE , FS , and FW respectively. Expressions for the Hazard Index in the four directions computed here are given as follows:

$$FN = (V/g)dv/dy - w/V$$

$$FS = -(V/g)dv/dy - w/V$$

$$FE = (V/g)du/dx - w/V$$

$$FW = -(V/g)du/dx - w/V$$

where u and v are the synthesized horizontal winds along the x (east) and y (north) axes respectively, w is the synthesized vertical air motion (positive upward), and V is the specified aircraft airspeed (140 kt). Horizontal gradients of the wind were computed with a three point central difference method applied to wind fields synthesized from the 31 May dual-Doppler radar measurements (Kropfli, 1986).

The above four Hazard Indices were computed for the four lowest horizontal levels of each of the eight processed radar volumes. These levels are at 75 m, 225 m, 375 m, and 525 m above ground. The ASCII formatted tapes and the contours provided in Appendix II include the fields FN , FS , FE , and FW along with all necessary housekeeping information. Table 1 indicates the start times (MST) of each of the eight volumes that were processed.

Table 1. Volume start times

Volume Number	Time of Start (MST)
1	1247:23
2	1249:03
3	1250:43
4	1252:23
5	1254:04
6	1255:44
7	1257:24
8	1259:04

Table 2 below lists the peak indices for all eight volumes as a function of height. It is apparent from the table that indices for this microburst would have been hazardous to aircraft approaching or taking off in some directions since it produced values that often exceeded .10, especially for the indices FE and FN. Values exceeding .12 were observed for FN near the surface for volumes 5 and 6 and also for FE, though only at altitudes of 225 m and above. The largest value observed was .15 for volume 6 at the lowest level for an approach or departure toward the north.

This microburst exhibits a strong asymmetry, with flight paths toward the east and the north having peak values nearly three times greater than values observed for the other directions. Flight paths through this microburst toward the north or the east would clearly be much more hazardous than flights toward the west or south.

The height dependence of the maxima varied differently with flight direction; FE peak values were generally increasing with height and FN was more nearly constant. The temporal behavior of the peak hazard index was somewhat erratic but generally the highest values were found in the three scans centered around 1254:04, as would be expected from a perusal of the synthesized wind fields (Kropfli, 1986).

The magnetic tape provided to NASA/Langley is card image format with 80 ASCII characters/record. Along with the data discussed above, the tape contains important housekeeping information such as the grid specifications, field names, radar coordinates, and volume start and end times. A description of the tape format is also provided in Appendix III along with the housekeeping information written to tape for all eight volumes.

Table 2. Peak hazard index as a function of time, direction and height

Index	Volume Number	Height (M AGL)			
		75	225	375	525
FN	1	.081	.063	.057	.070
	2	.084	.063	.065	.079
	3	.091	.087	.084	.067
	4	.098	.098	.128	.134
	5	.120	.119	.105	.097
	6	.152	.118	.096	.099
	7	.076	.083	.089	.103
	8	.071	.061	.070	.068
FS	1	.055	.033	.039	.055
	2	.053	.034	.045	.056
	3	.053	.044	.061	.089
	4	.049	.049	.063	.079
	5	.065	.072	.078	.113
	6	.065	.059	.092	.105
	7	.057	.070	.081	.100
	8	.050	.055	.068	.085
FE	1	.057	.054	.058	.057
	2	.075	.065	.076	.102
	3	.071	.096	.115	.121
	4	.096	.110	.128	.149
	5	.082	.121	.140	.141
	6	.095	.091	.113	.137
	7	.095	.086	.115	.129
	8	.070	.089	.100	.092
FW	1	.040	.038	.039	.046
	2	.040	.035	.044	.050
	3	.039	.049	.045	.064
	4	.048	.053	.053	.059
	5	.039	.044	.059	.068
	6	.058	.060	.075	.057
	7	.057	.048	.046	.058
	8	.048	.041	.043	.055

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APPENDIX I

Reprint of: A Microburst Observed by High-Resolution Dual-Doppler Radar

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A MICROBURST OBSERVED BY HIGH-RESOLUTION DUAL-DOPPLER RADAR

R. A. Kropfli
NOAA/ERL/Wave Propagation Laboratory
Boulder, Colorado 80303

1. INTRODUCTION

The Phoenix II convective boundary layer (CBL) experiment was conducted near the Boulder Atmospheric Observatory (BAO) instrumented tower during the summer of 1984. The experiment, a collaborative effort between the University of Oklahoma, National Oceanic and Atmospheric Administration (NOAA), and National Center for Atmospheric Research (NCAR), addressed the small-scale structure of the CBL with Doppler radar, instrumented aircraft, and surface measurements. Scan procedures and the 13 km radar separation optimized measurements of the small-scale and short-lived eddies present in the dry CBL. This experimental design was also suitable for documenting microbursts, which often develop and decay within a 10 min period and have downdrafts that are as small as 1 km in diameter just above the surface (Wilson et al., 1984; Fujita, 1981). Two adjacent areas of dual-Doppler radar coverage were established: one formed by the two NCAR C-band radars and the other formed by the two NOAA/WPL X-band radars. The 31 May microburst discussed here was centrally located within the scan area of the NOAA radars, and these two radars are used in the analysis.

There were three confirmed reports of microbursts on this day during which virga and many small rainshafts were observed; it is suspected that a large number of such events were occurring throughout the region. Confirmation of one of these events came from a jetliner departing from Stapleton Airport about 30 km to the southeast of the Phoenix II measurement area. As a result of a 10 m s^{-1} loss of airspeed during acceleration down the runway, the aircraft failed to clear a 5-m-high radio beacon located 300 m beyond the end of the runway resulting in major damage to the fuselage. The loss of airspeed was later determined to be the result of a moderately strong microburst that impacted the runway at around 1235 MST (NTSB, 1985). Fortunately, the aircraft climbed out of danger and returned safely to Stapleton with no injury or loss of life.

The event documented here took place about 20 minutes later near the BAO, and a third microburst was evident in the anemometer records at the BAO several hours later (Bedard, personal communication). This near-tragedy at Stapleton along with other well-documented aircraft accidents (Fujita and Byers, 1977) underscores the importance of finding reliable precursor signatures in Doppler radar observations of microbursts. Several features of the 31 May microburst are apparent in the radar data that might be useful in identifying microbursts several minutes before the outflow reaches its maximum intensity at the surface.

2. ENVIRONMENTAL CONDITIONS

The 12 Z (0600 MST) Denver sounding shown in Fig. 1 suggests that the likelihood of microbursts occurring on this day was high. Note especially the high lapse rate (8.9 deg/km) just below the 500 mb level, the high relative humidity above 600 mb, and the relatively dry layer below 600 mb. These features have been associated with the occurrence of microbursts (Ciracena et al., 1983; Bedard and LeFebvre, 1986). The NCAR instrumented King Air aircraft had been making soundings within and near the area of dual-Doppler radar observations (Hildebrand, personal communication); the potential temperature and mixing ratio profiles from the sounding at 1236 MST are shown in Fig. 2. A well-mixed layer is seen below 2 km AGL, and slightly stable air is present from 2 km to 3 km AGL, the highest level flown. (All heights are AGL.) The melting level was found to be at 2.8 km and cloud base was at 3.8 km. Figure 3 shows the mean wind profile derived from the dual-Doppler analysis at 1250:45 and indicates moderate shear toward the northeast at about 0.003 s^{-1} . The mean profiles of wind, stability, and relative humidity were critical to the overall physical structure of this event.

3. PHYSICAL CHARACTERISTICS

Figure 4 indicates that the microburst near the BAO was in an ideal location relative to the two NOAA X-band radars. At the time of the incident at Stapleton, these two radars were being used in high resolution observations of the

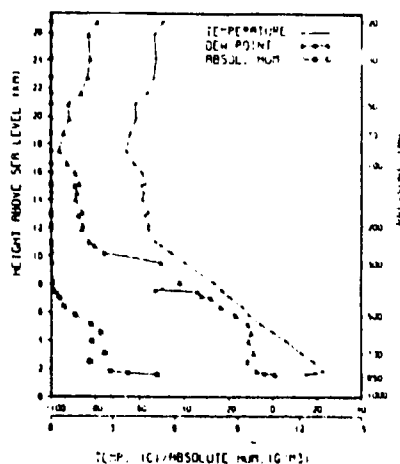


Figure 1. Temperature and humidity profiles from the MWS morning radiosonde at Denver on 31 May 1984.

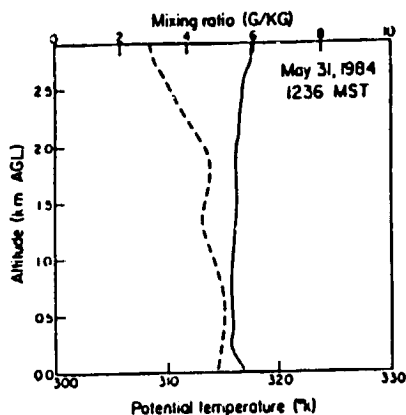


Figure 2. Potential temperature (solid) and mixing ratio (dashed) profiles from the NCAR King Air aircraft taken near the BAO at 1236 MST on 31 May 1984.

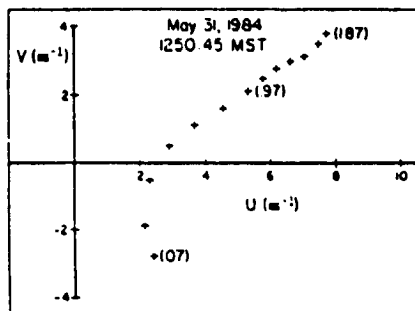


Figure 3. Hodograph obtained from the Doppler radar derived wind field at 1250:45 MST on 31 May 1984. Height AGL in km shown in parentheses.

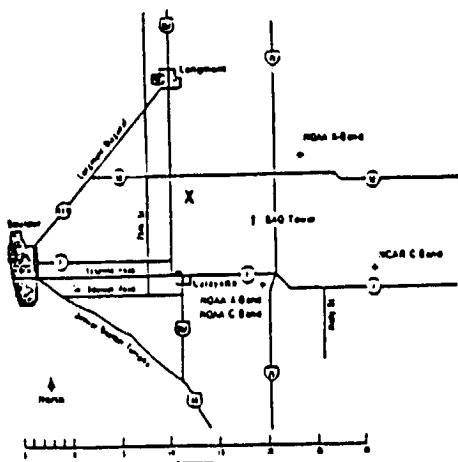


Figure 4. Phoenix II experimental layout showing approximate location of the microburst, indicated by the X.

CBL to a height of 2.5 km and with a scan cycle time of 100 s. These measurements permitted use of the 150 m grid spacing shown in Figs. 5 to 9. A 0.7 μ s pulse width and a beam separation of 0.7 deg in azimuth were used in these measurements.

Figure 5 shows the east-west cross section through the core of the microburst, the result of a synthesis of the dual-Doppler radar data taken at 1254:04. (In all the flow fields shown here the volume mean of approximately 5.2 m s^{-1} to the ENE was removed to display the flow features more clearly.) It is apparent that the downdraft has its source region above 2.5 km AGL . A maximum downward air motion of nearly 8 m s^{-1} occurs at about 700 m .

The sequence of horizontal views at $z = 75$ m in Fig. 6 shows the rapid development of the divergent outflow near the surface at 100 s intervals. Figure 6c shows the surface outflow at maximum intensity when a peak divergence of 0.02 s^{-1} was observed. The maximum surface reflectivity of 39 dBZ occurred 5 min earlier at 1250 and decreased gradually to 36 dBZ by 1256.

A comparison of Fig. 7, showing the horizontal flow at 375 m, with the low level flow at the same time (1252:23) in Fig. 6a suggests that the divergent outflow was readily observable above the surface nearly 200 s before it reached its maximum at the surface. In this case horizontal divergence at a height of several hundred meters might have been a precursor of strong, potentially dangerous outflows at the surface. It is also conceivable that a single radar, appropriately located relative to the microburst might have detected this precursor to the surface outflow.

A second precursor, midlevel convergence, is suggested by Fig. 8. These data were taken at 1250:43 and at a height of nearly 1.9 km. It is

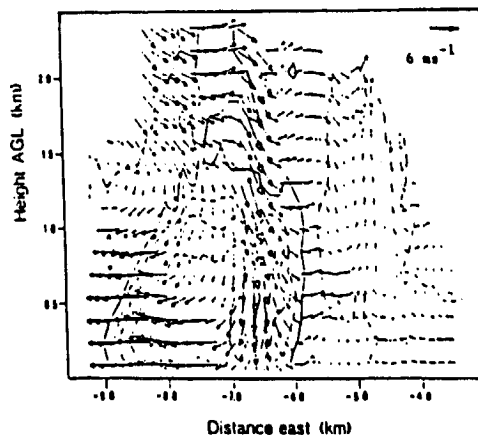


Figure 5. East-west cross-section of dual-Doppler eddy wind field through core of the microburst at 1254:04 MST. The cross section is taken 3.0 km north of the BAO. The solid contour represents the 9.5 dBZ reflectivity, and the dashed lines are at 5.5 dB increments.

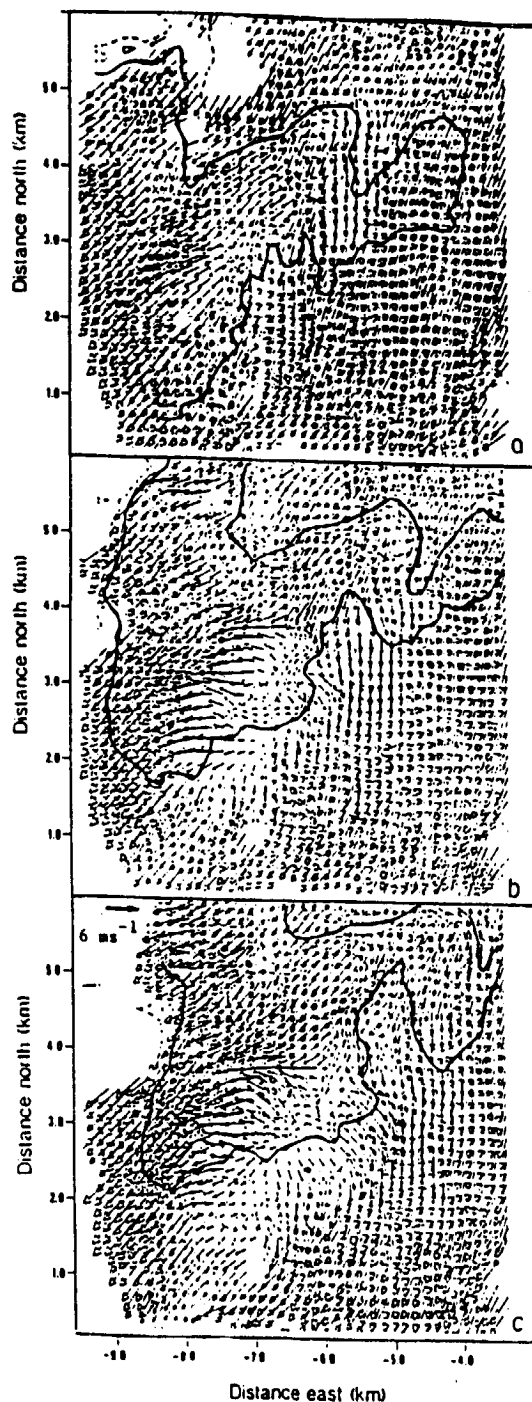


Figure 6. The horizontal eddy wind field at $z = 75$ m. Arrows are scaled as in Fig. 5 and reflectivity contours are approximately the same as in Fig. 5. (a) 1252:23 MST; (b) 1254:04 MST; (c) 1255:44 MST.

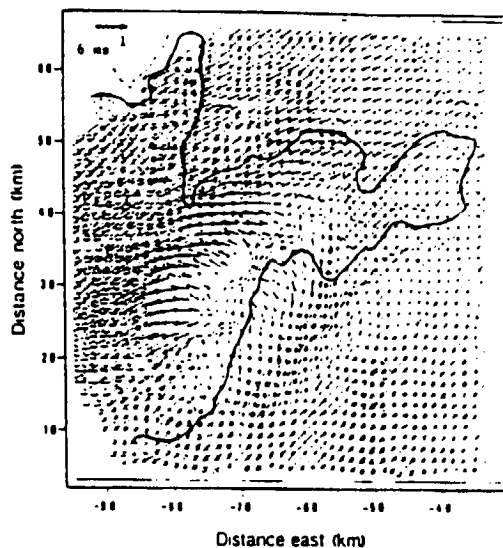


Figure 7. As in Fig. 6a except at $z = 375$ m.

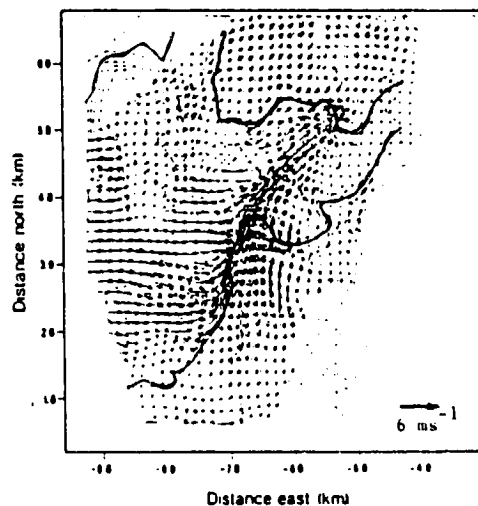


Figure 8. As in Fig. 6a except at $z = 1.9$ km and at 1250:43 MST.

likely that air near the 2 km level was entrained into the downdraft that had been initiated by precipitation from the cloud above 4.8 km. Convergence near the bottom of a slightly stable layer may indicate that strong outflows will soon be found at the surface. In this case the convergence at 2 km preceded the maximum divergence at the surface by nearly 5 min. Table 1 summarizes some of the properties of the microburst during its time of maximum outflow at 75 m.

An unexpected feature of these observations is the movement of the reflectivity maximum away from the peak downdraft region near the surface. Figure 9 indicates this with every other scan from 1250:43 to 1257:24 being shown; i.e., the samples in the figure are at 200 s intervals. The regions enclosed by solid con-

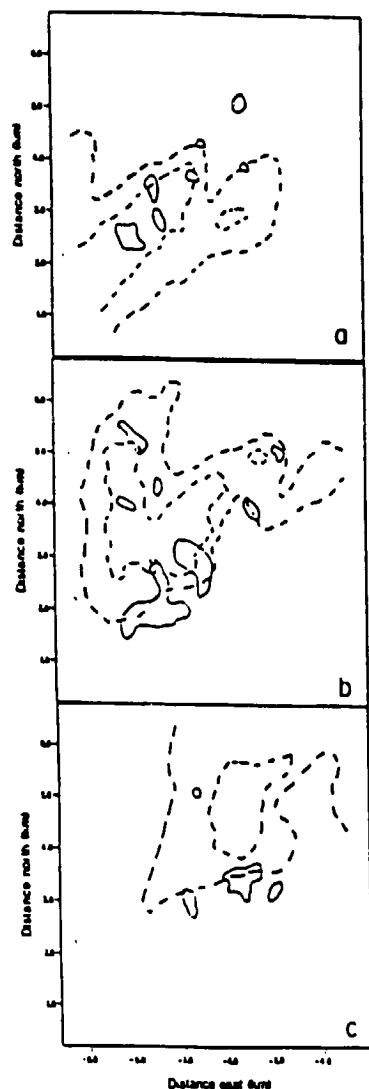


Figure 9. Reflectivity contours at 10 dBZ (dashed) and 24 dBZ (dashed) and $w = -1.0 \text{ m s}^{-1}$ (solid) at 1250:43 MST and at $z = 75 \text{ m}$. (a) 1250:43 MST; (b) 1254:04 MST; (c) 1257:24 MST.

Table 1. Summary of microburst properties at (1255:44 MST) time of maximum low-level divergence

Peak divergence at 75 m:	$2.1 \times 10^{-2} \text{ s}^{-1}$
Peak reflectivity at 75 m:	36 dBZ
Peak downdraft and height:	6.5 m s^{-1} , .8 km AGL
Area with $\text{DIV} > 10^{-2} \text{ s}^{-1}$ at 75 m:	.4 km^2
Mean flow in sample volume:	5.2 m s^{-1} from WSW
Time after first appearance of convergence at ~2 km:	~5 min
Time after first appearance of divergence at ~.4 km:	~3 min

tours indicate $w < -1 \text{ m s}^{-1}$ and the inner dashed contour encloses reflectivity greater than 24 dBZ. w is the vertical air motion. The sequence clearly shows the motion of the heavier precipitation moving to the northeast while the main downdraft region remains nearly stationary. (The three minor downdrafts in the northwest corner of Fig. 9b appear only briefly and are not thought to be of significance.) Near the end of the observations at 1257:24 the decaying downdraft is located almost entirely outside of the 10 dBZ contour.

The preliminary explanation offered for this unexpected observation is that the downdraft is initiated by light rain, in which case water loading could play a significant role along with the negative buoyancy provided by melting and evaporative cooling. With time, melting and evaporative cooling becomes more dominant on the upshear (southwest) side of the precipitation shaft. In effect the source region for the downdraft moves upshear relative to the reflectivity core and into a region of smaller hydrometeors which may be more likely to provide stronger downdrafts as a result of enhanced evaporative cooling (Srivastava, 1985). The reasons for this upshear movement of the source region are not understood, and, unfortunately, the radar observations were not high enough to provide additional insight.

5. SUMMARY

The high resolution Doppler radar observations described here suggest two possible precursor signatures to strong surface outflows associated with microbursts: divergence a few hundred meters above the surface and convergence at heights several kilometers above the surface. We recognize that these links have not been seen before and that the observation of these features prior to surface outflow in this one case merely suggests that future high quality data sets should be examined with this in mind. The separation of the downdraft from the reflectivity core with the passage of time suggests that melting and evaporative cooling play an increasingly dominant role after precipitation loading initiates the downdraft.

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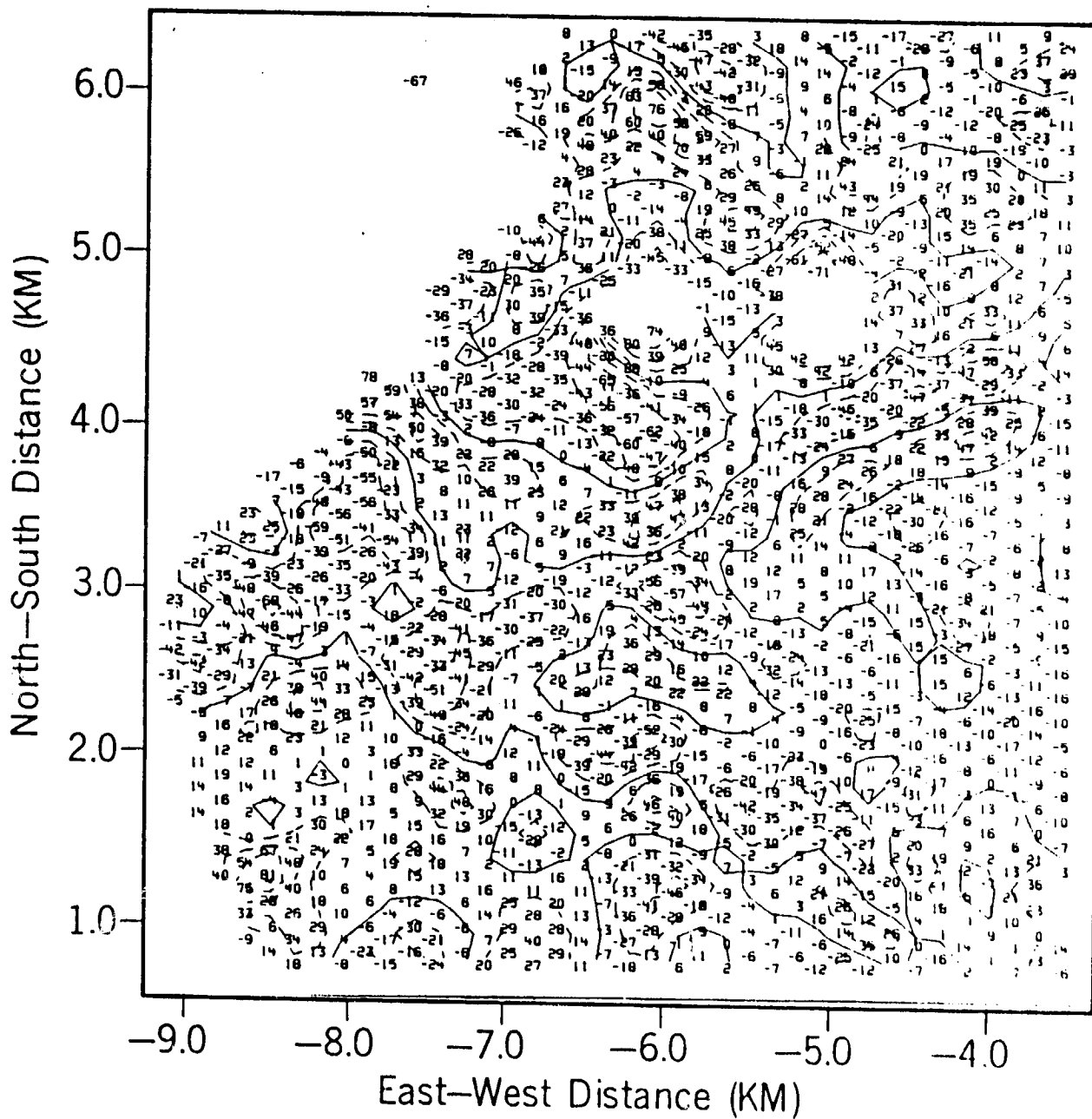
APPENDIX II

Hazard Index Plots

$Z = 75 \text{ m AGL}$

Field: FN X 1000

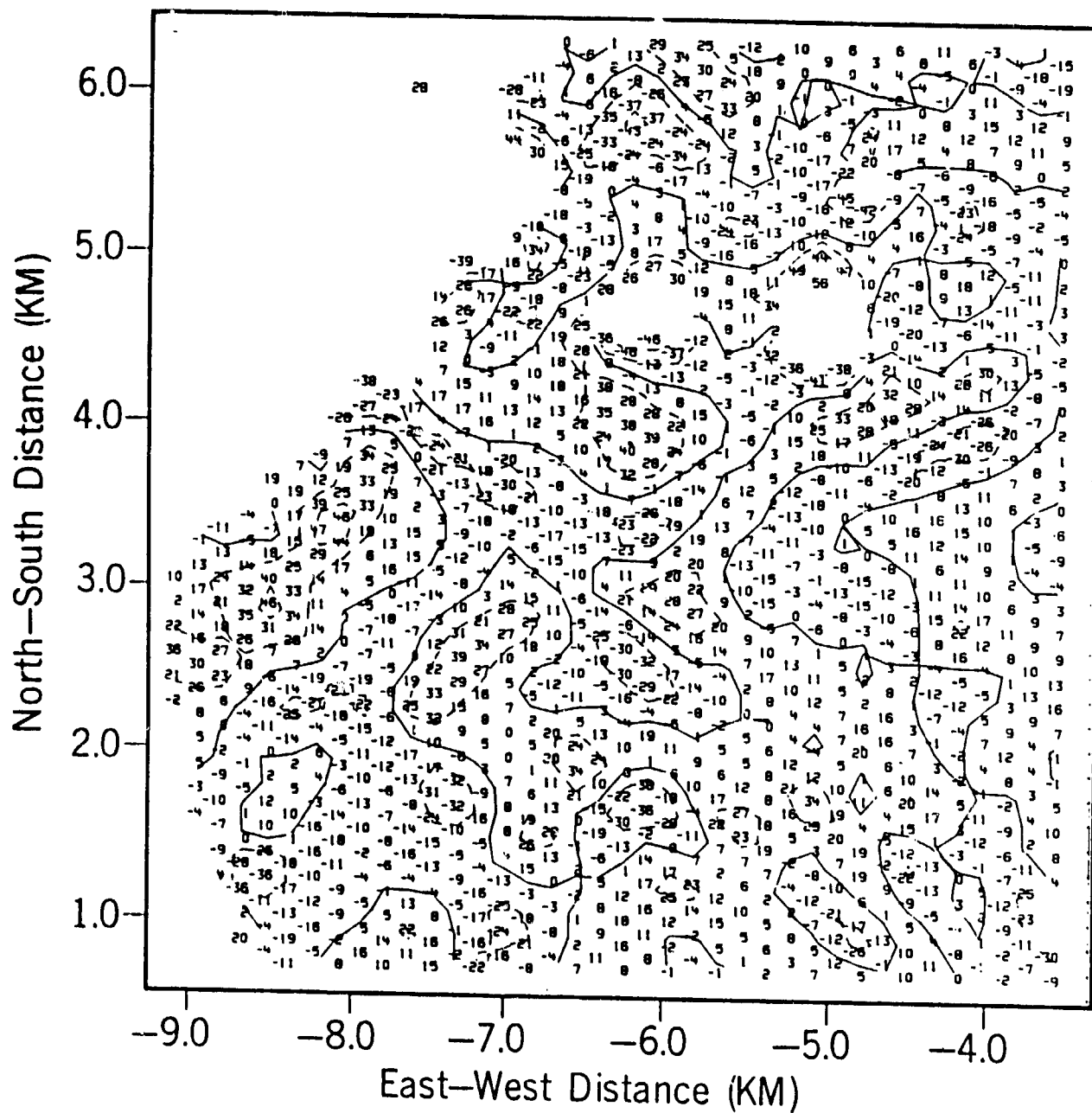
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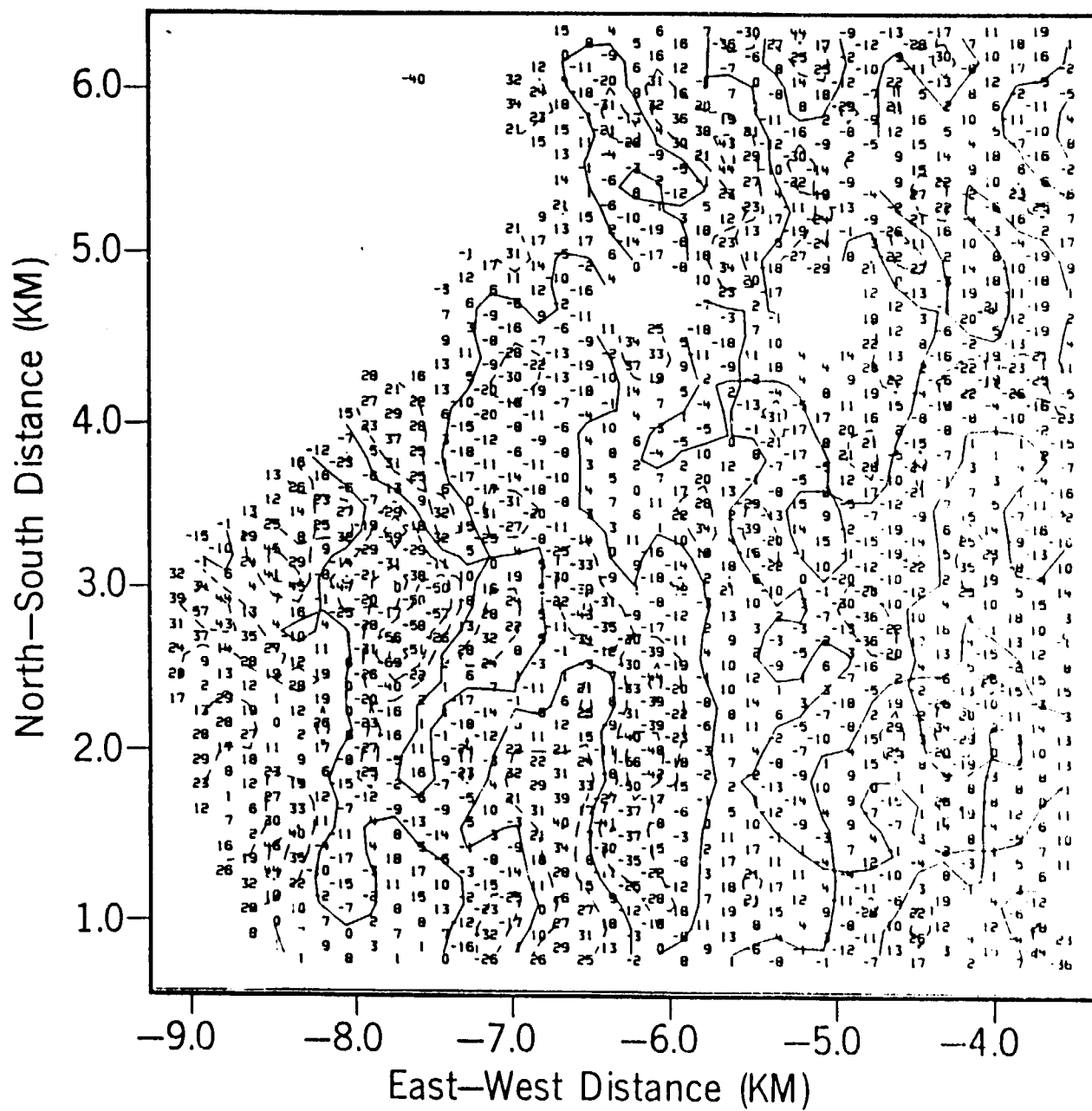
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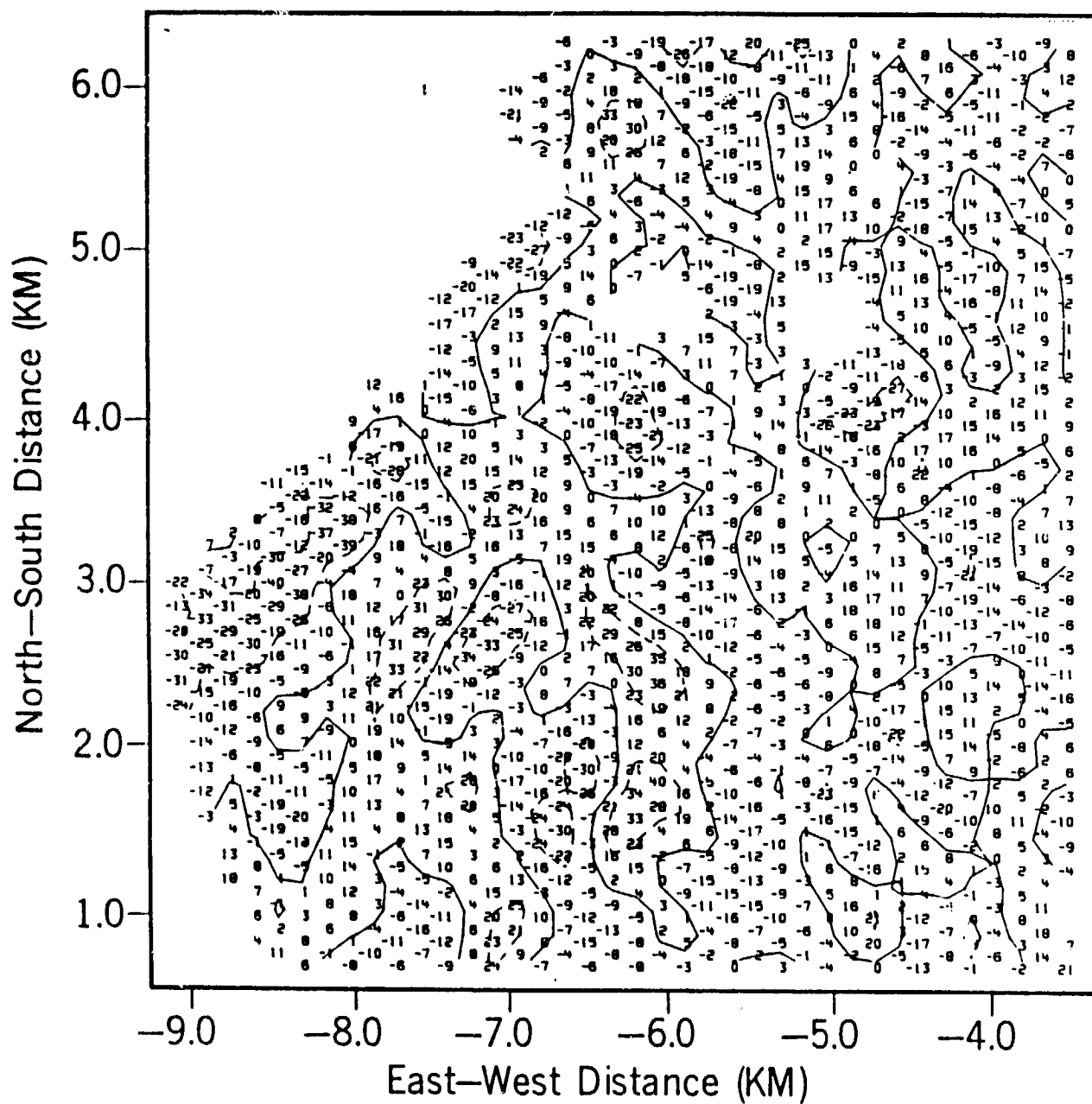
Field: FE X 1000

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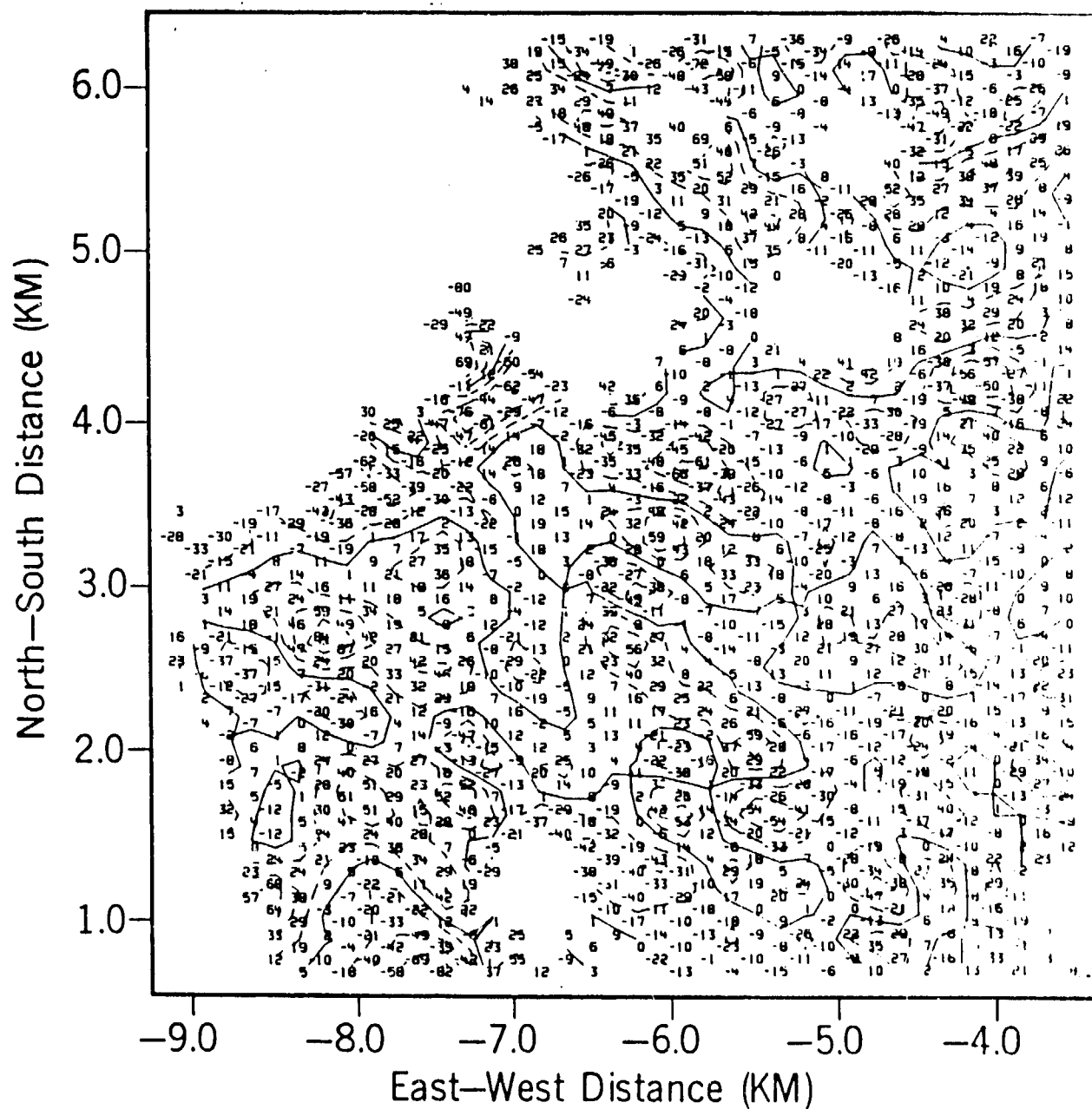
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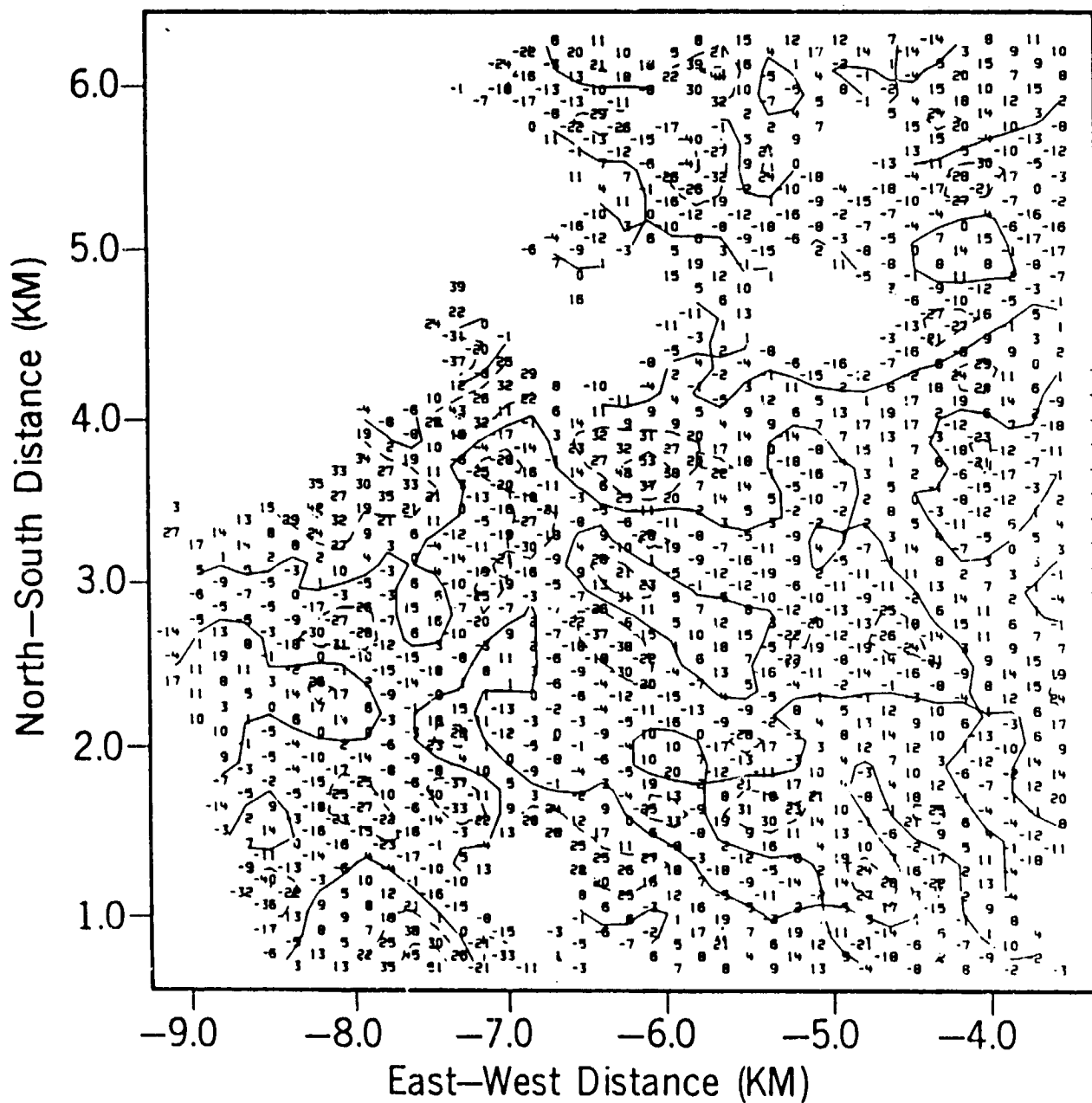
Field: [N X 1000

Time 1249:03



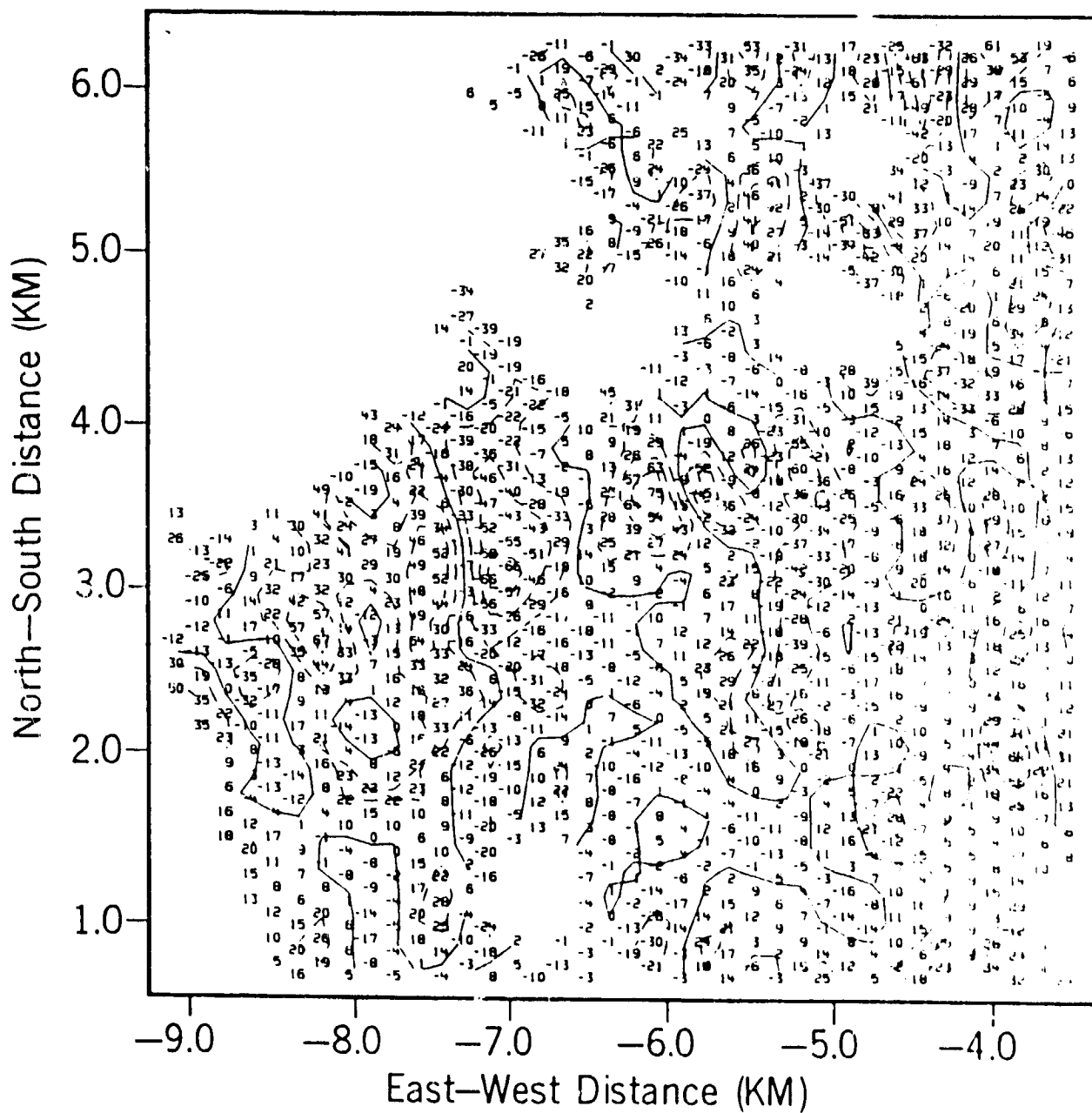
Field: FS X 1000

Time 1249:03



Field: FE X 1000

Time 1249:03

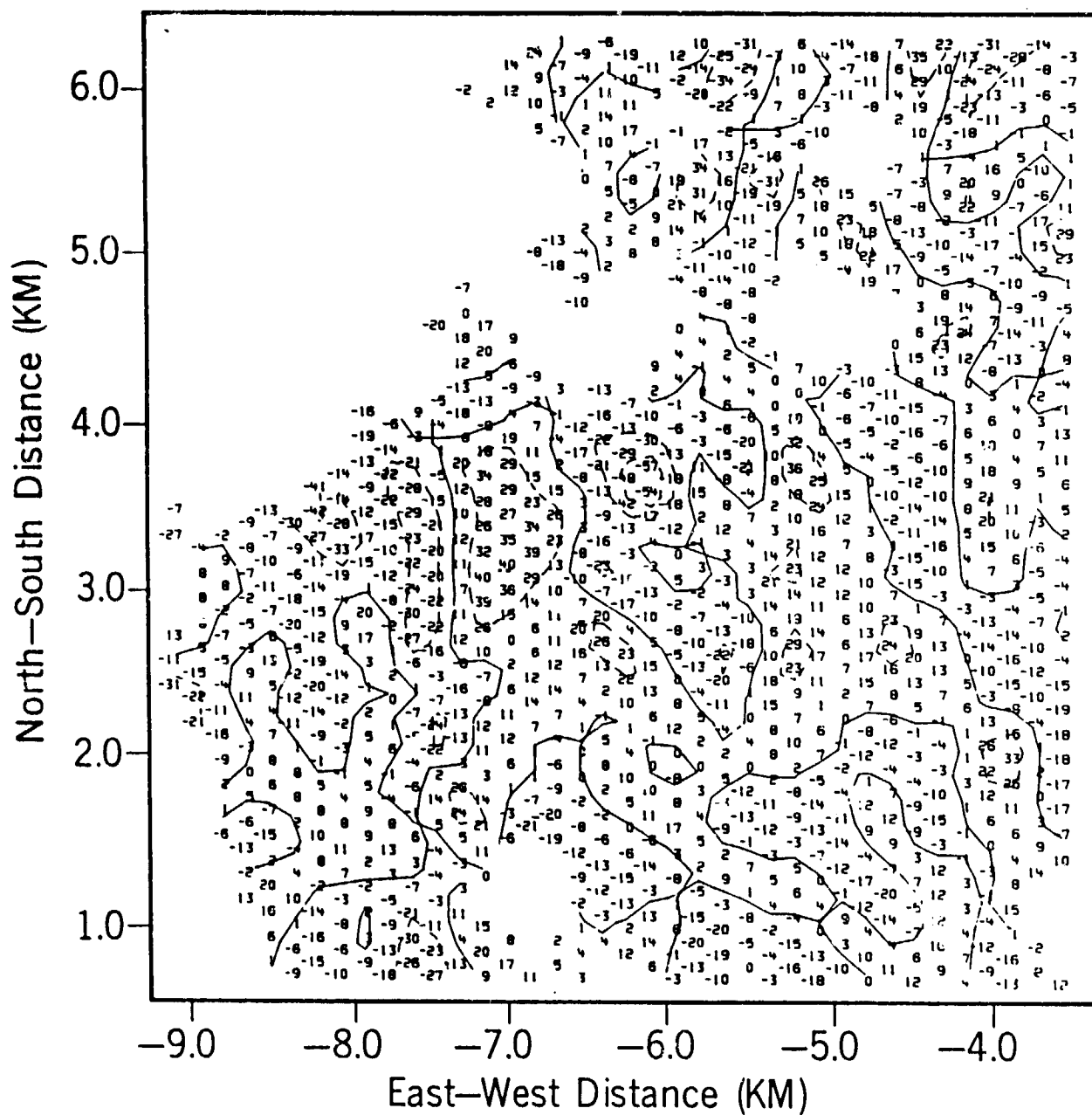


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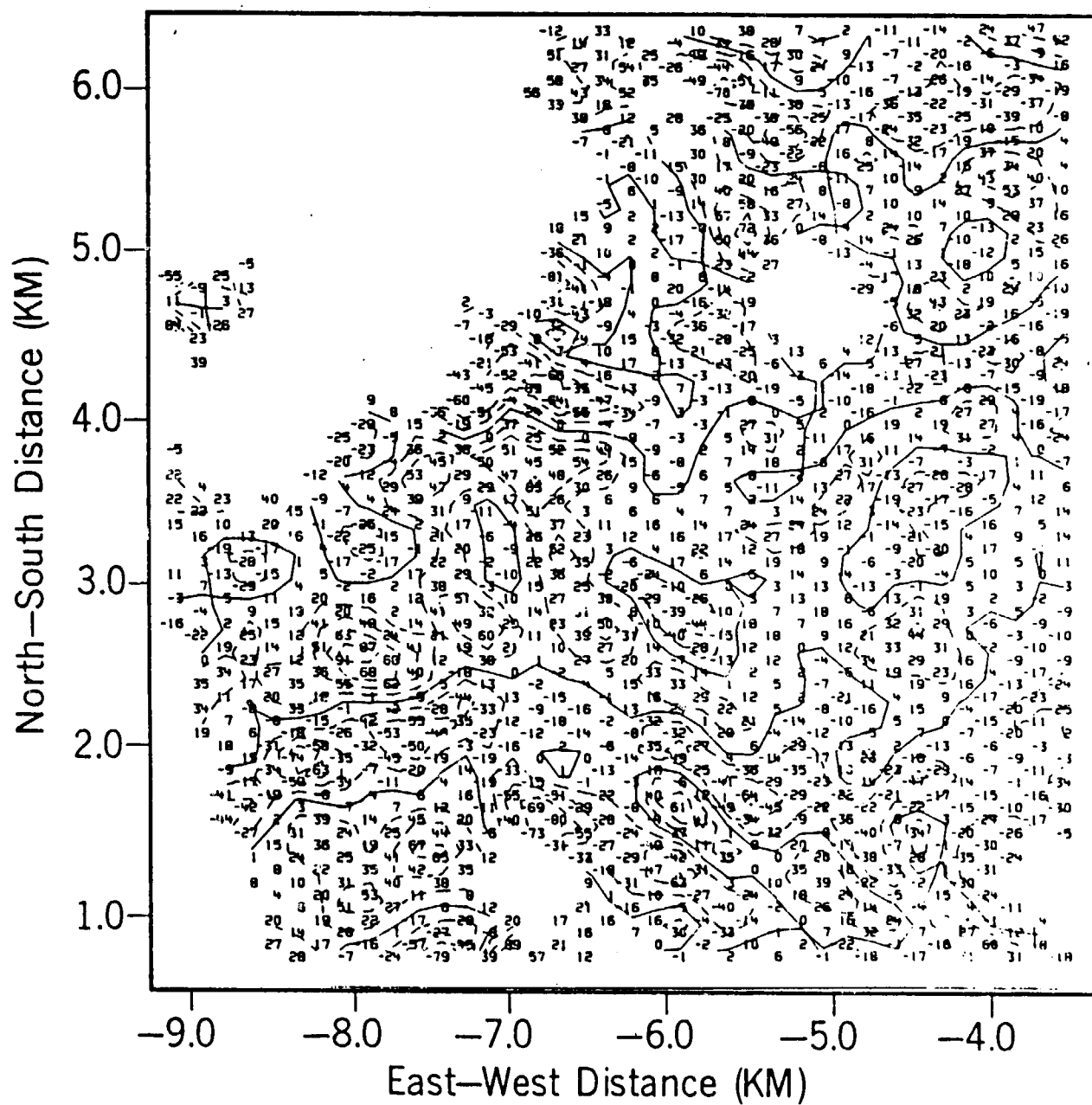
Field: FW X 1000

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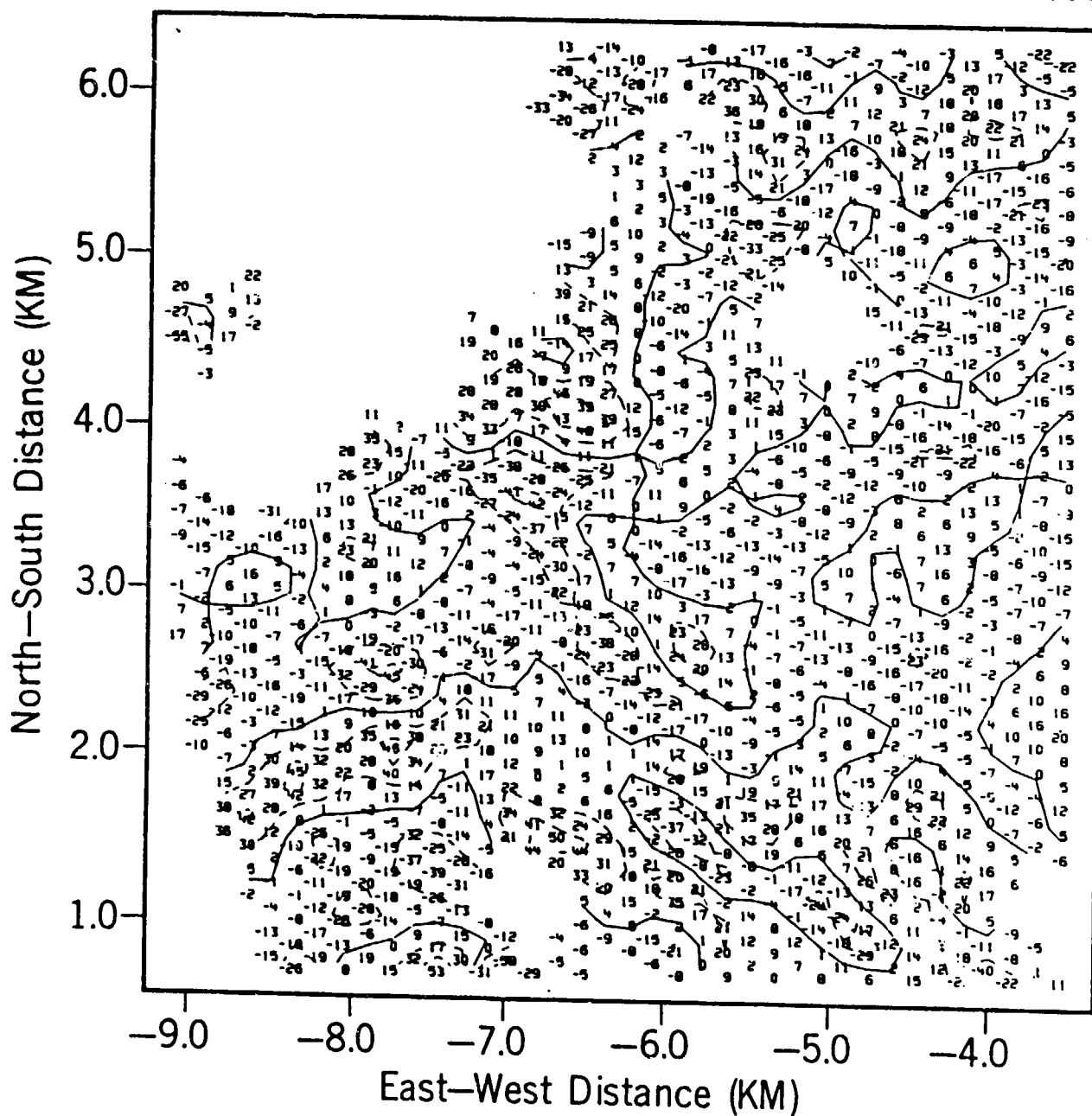
Field: FN X 1000

Time 1250:43



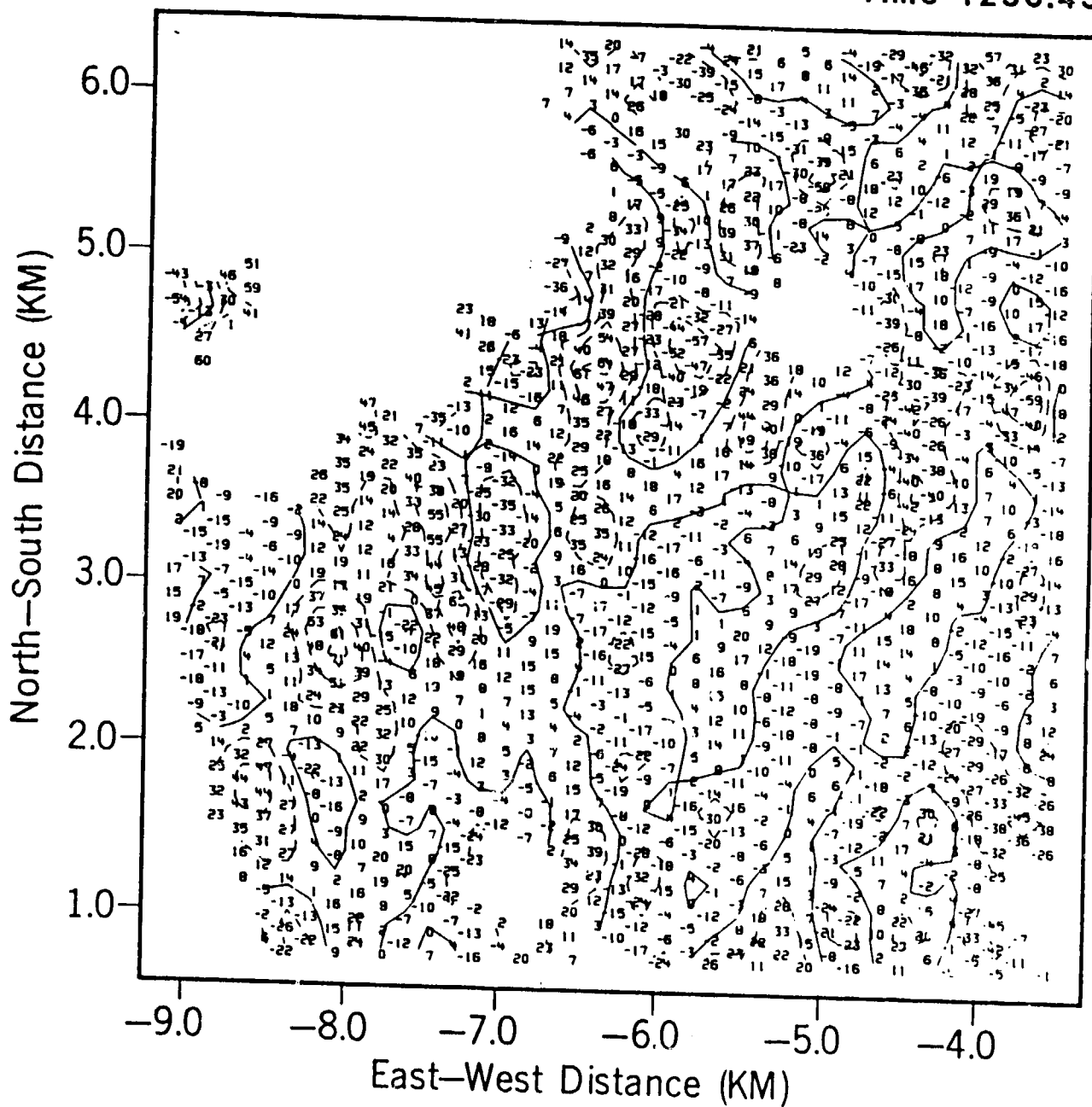
Field: FS X 1000

Time 1250:43



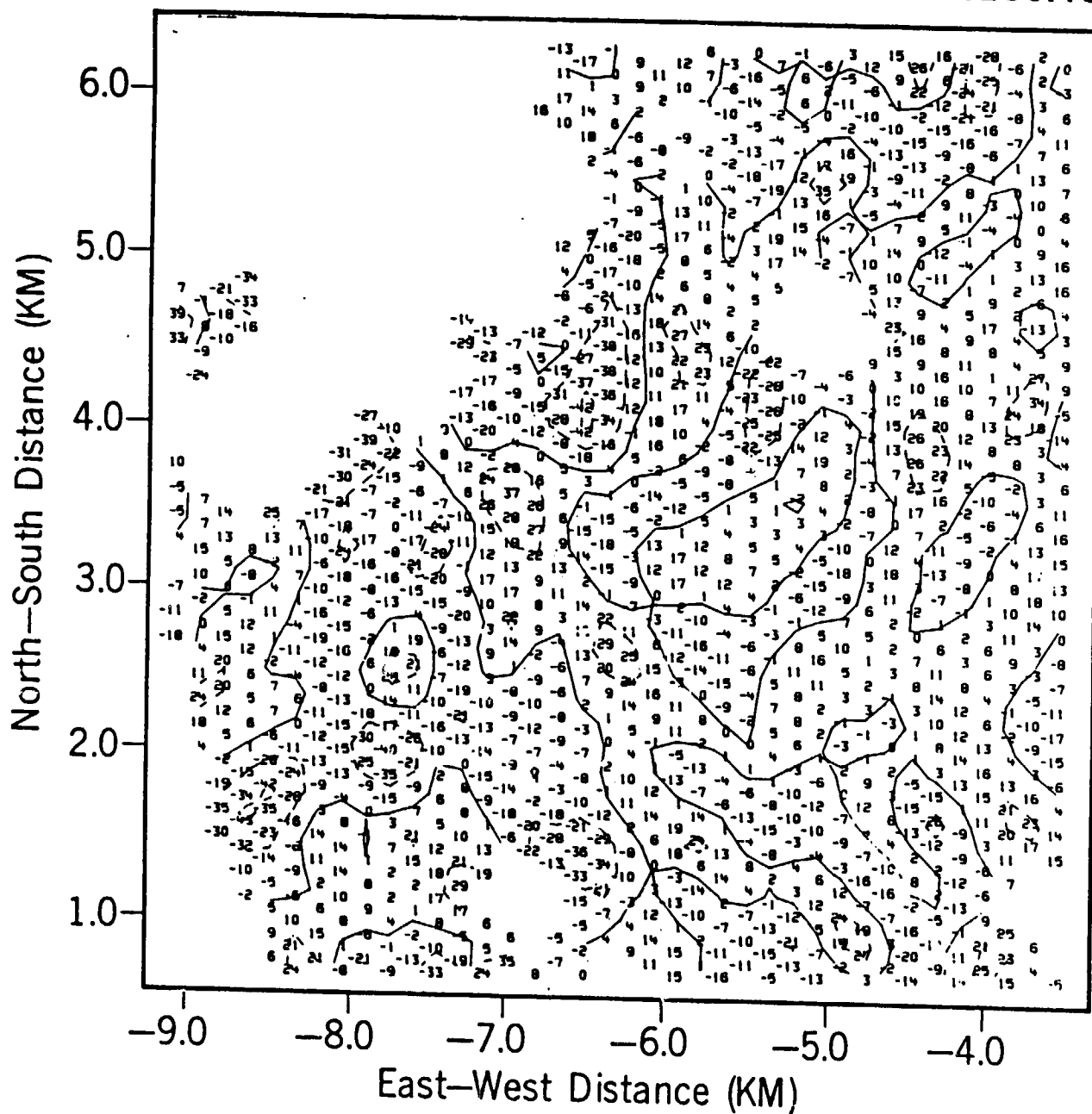
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Time 1250:43

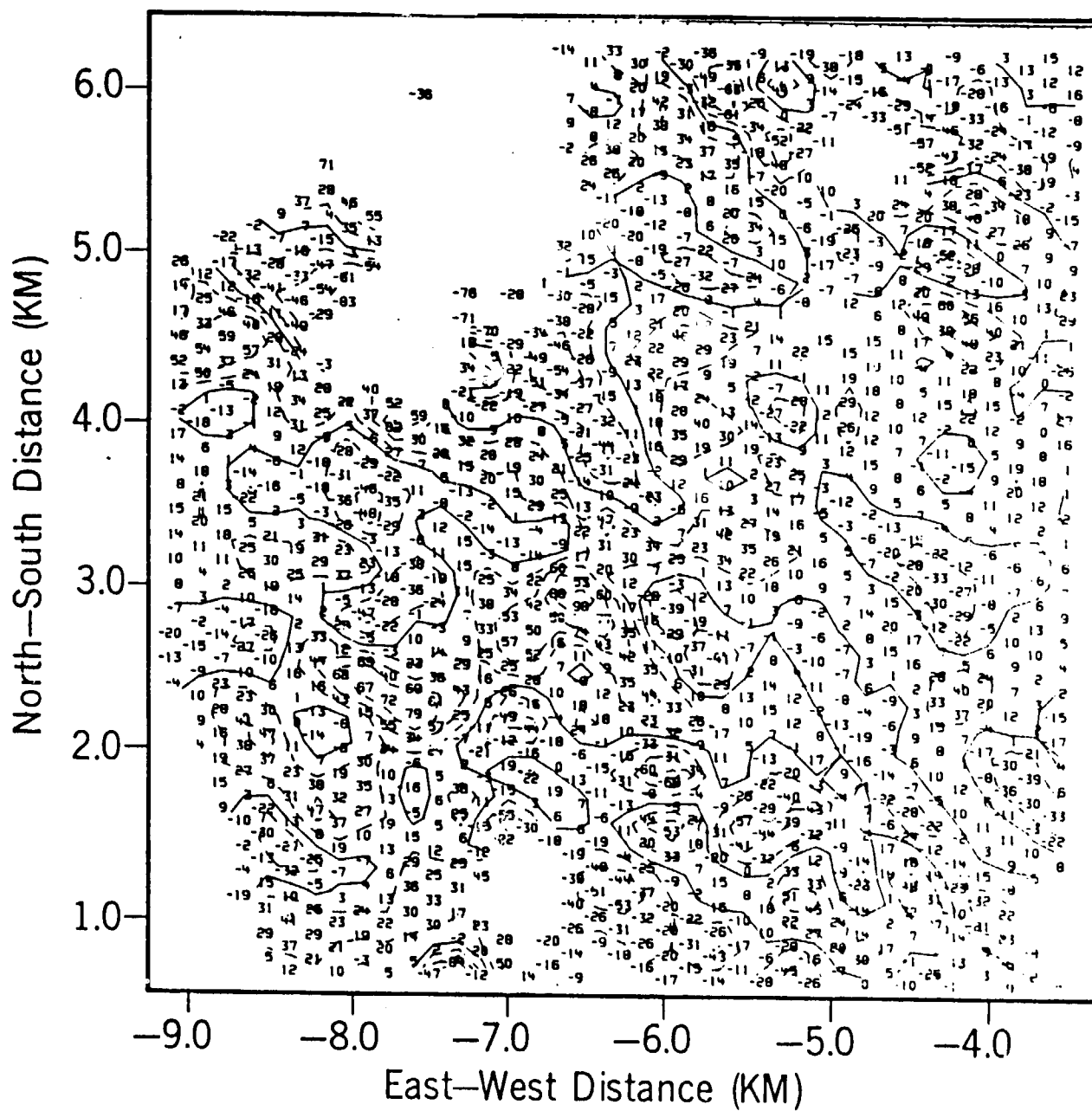


Field: FW X 1000

Time 1250:43

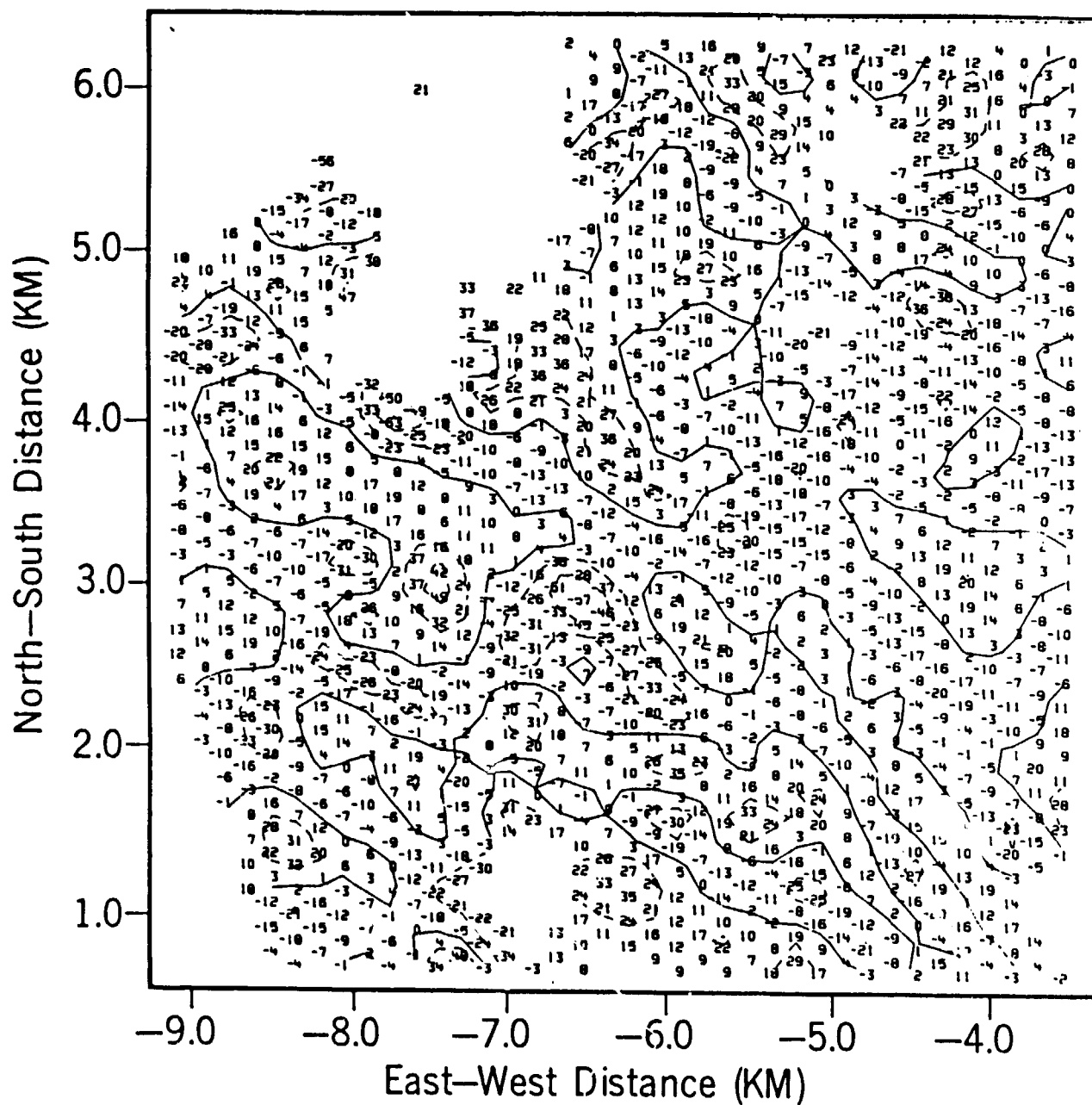


Time 1252:23



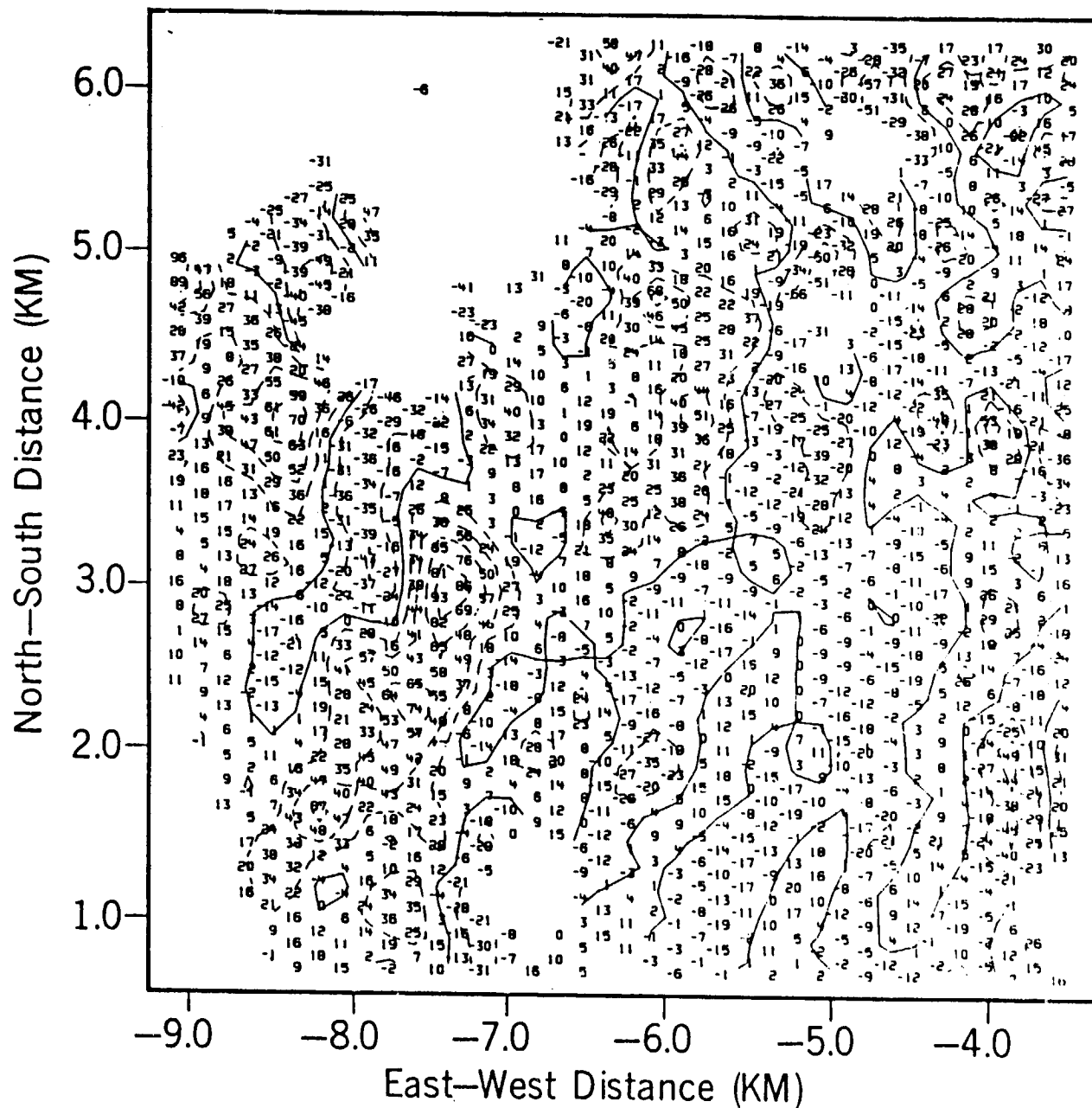
Field: FS X 1000

Time 1252:23



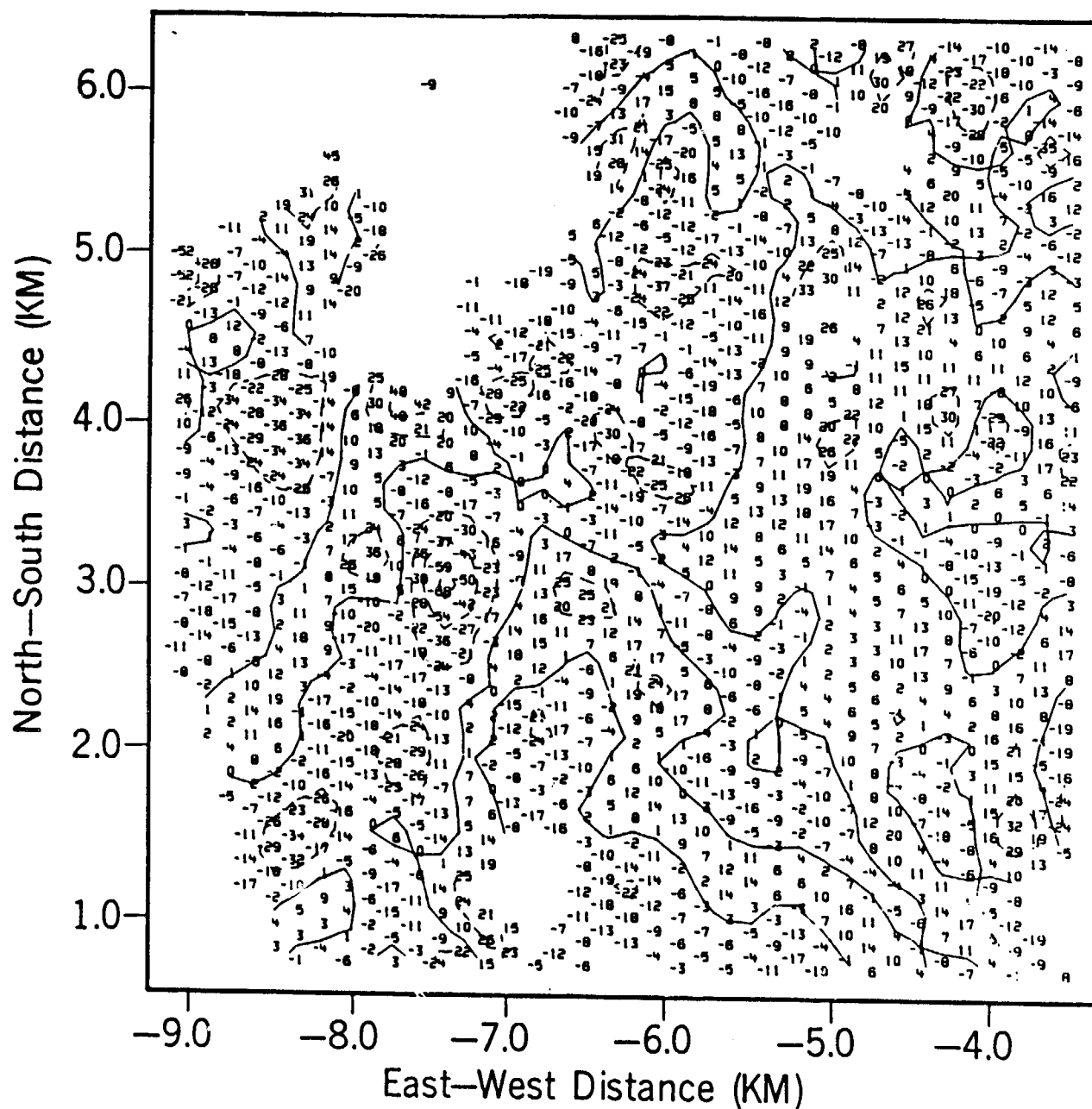
Field: FE X 1000

Time 1252:23



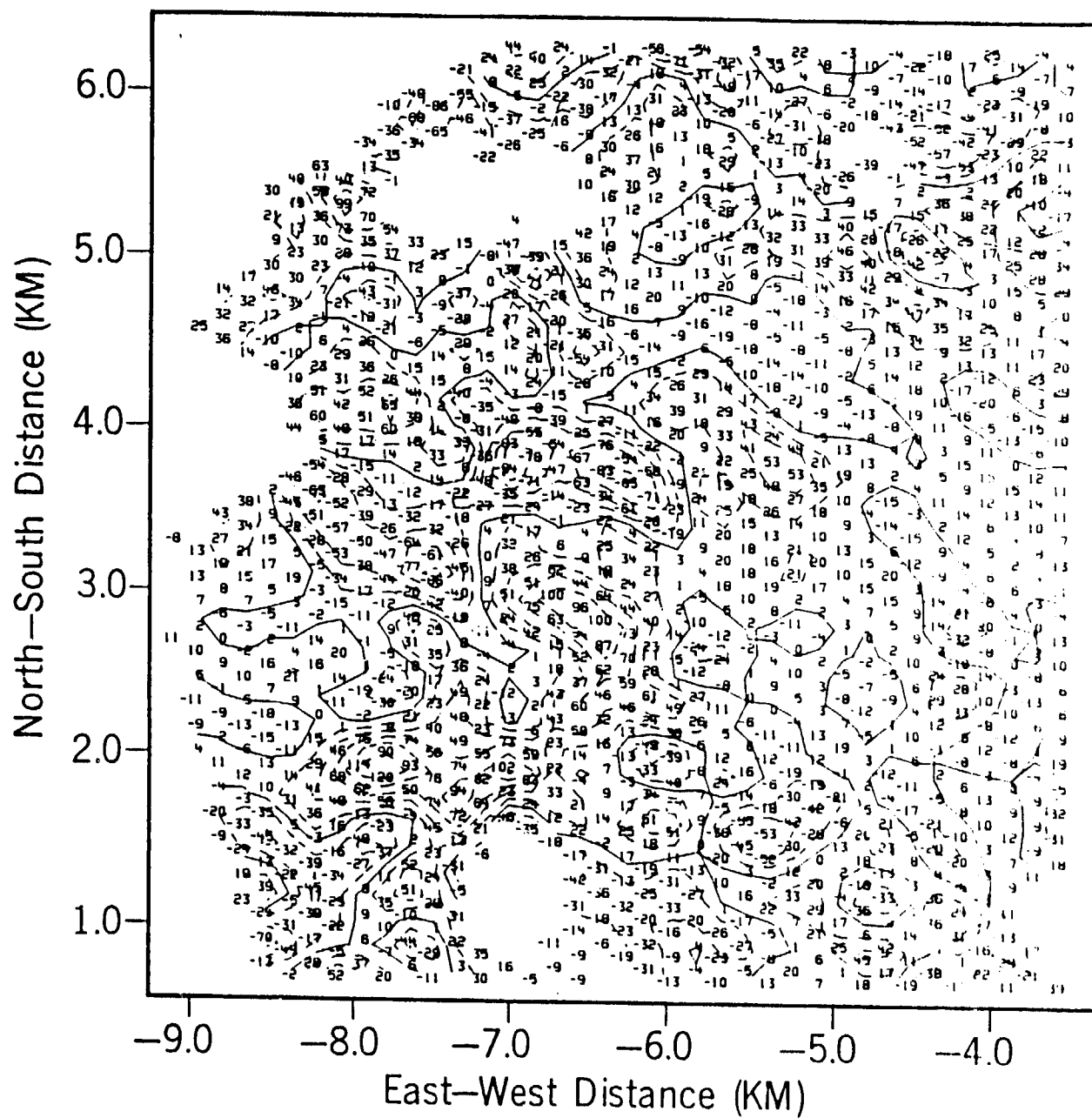
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Time 1252:23



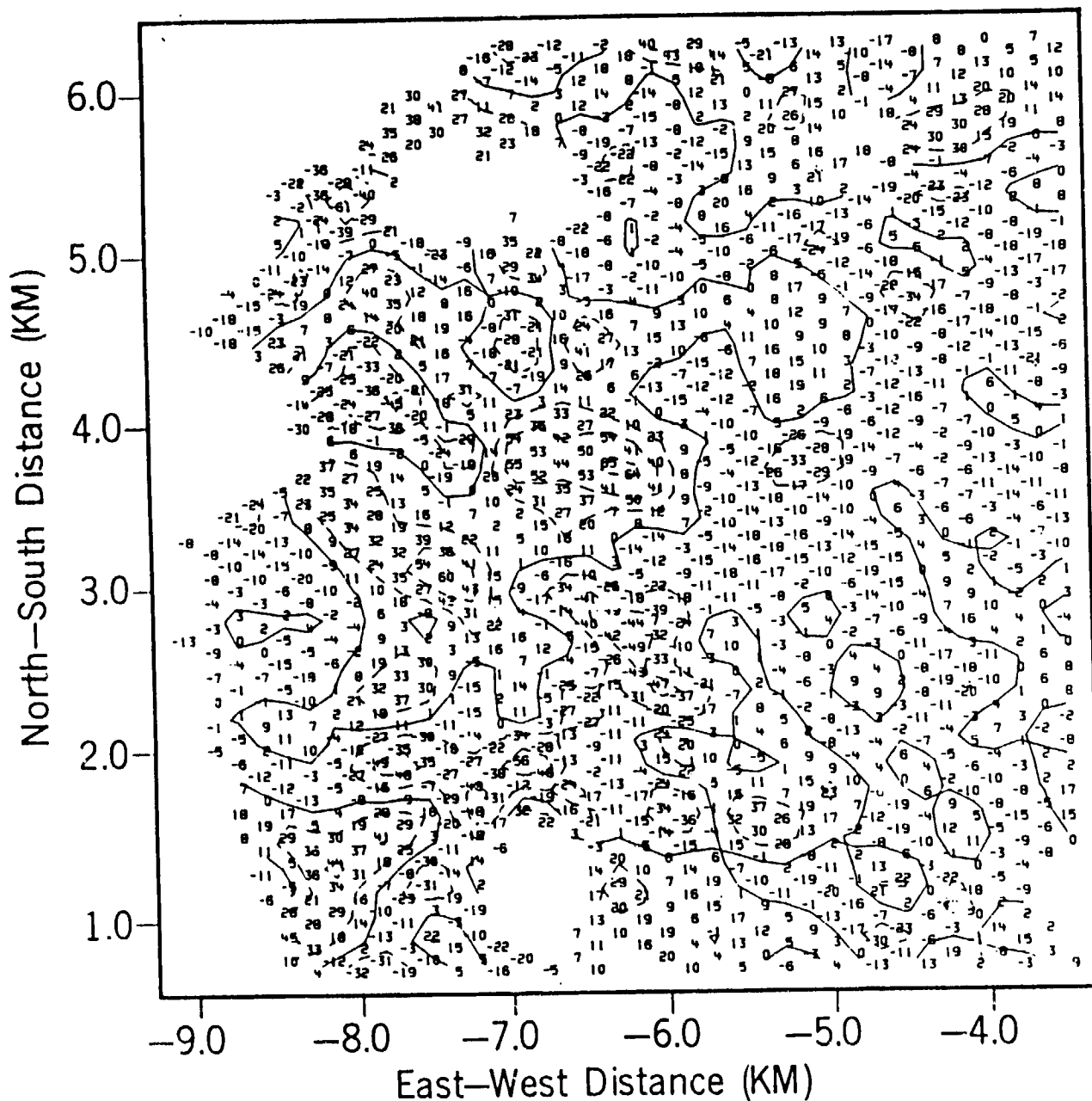
Field: FN X 1000

Time 1254:04



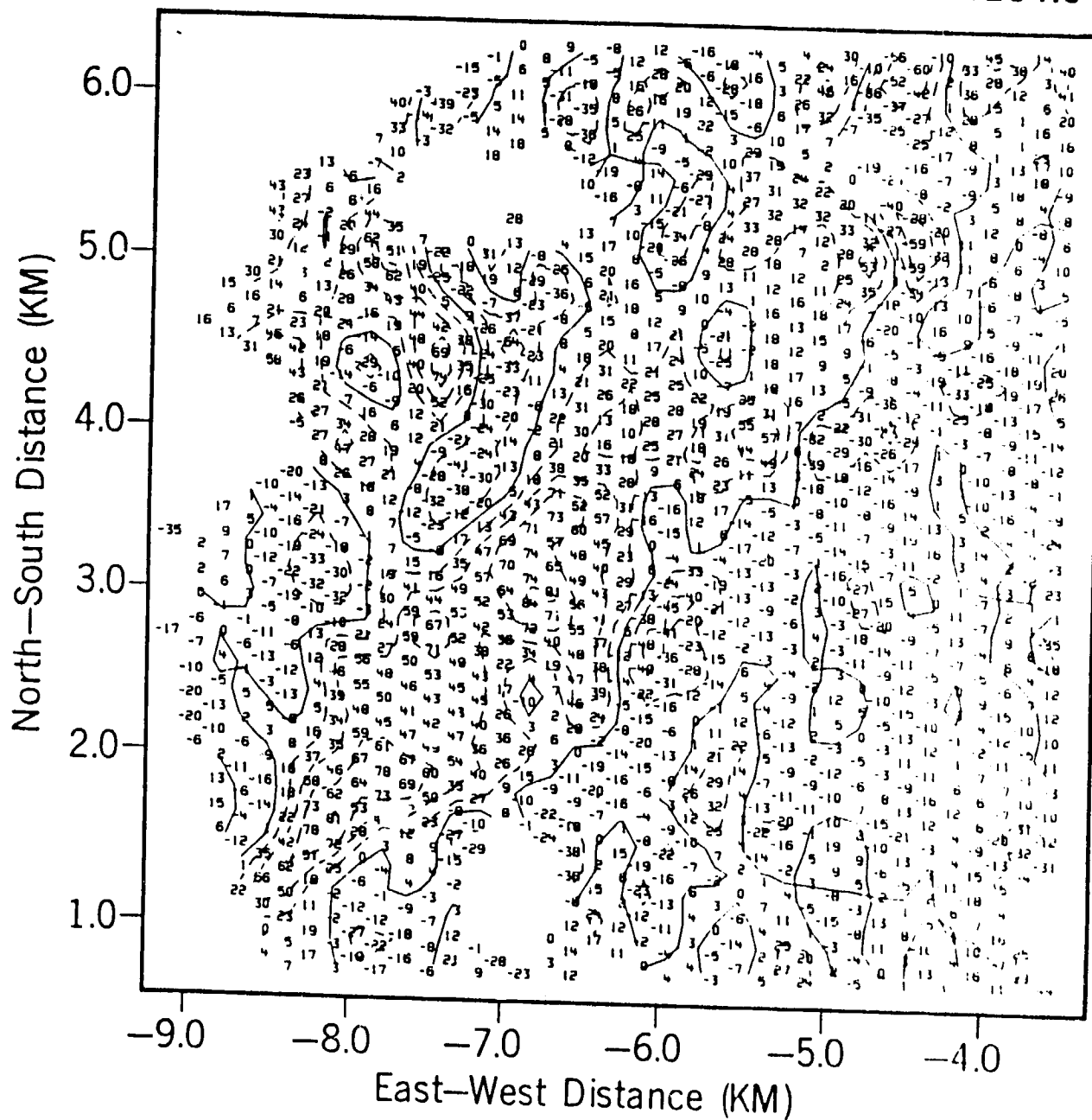
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Time 1254:04



Field: FE X 1000

Time 1254:04

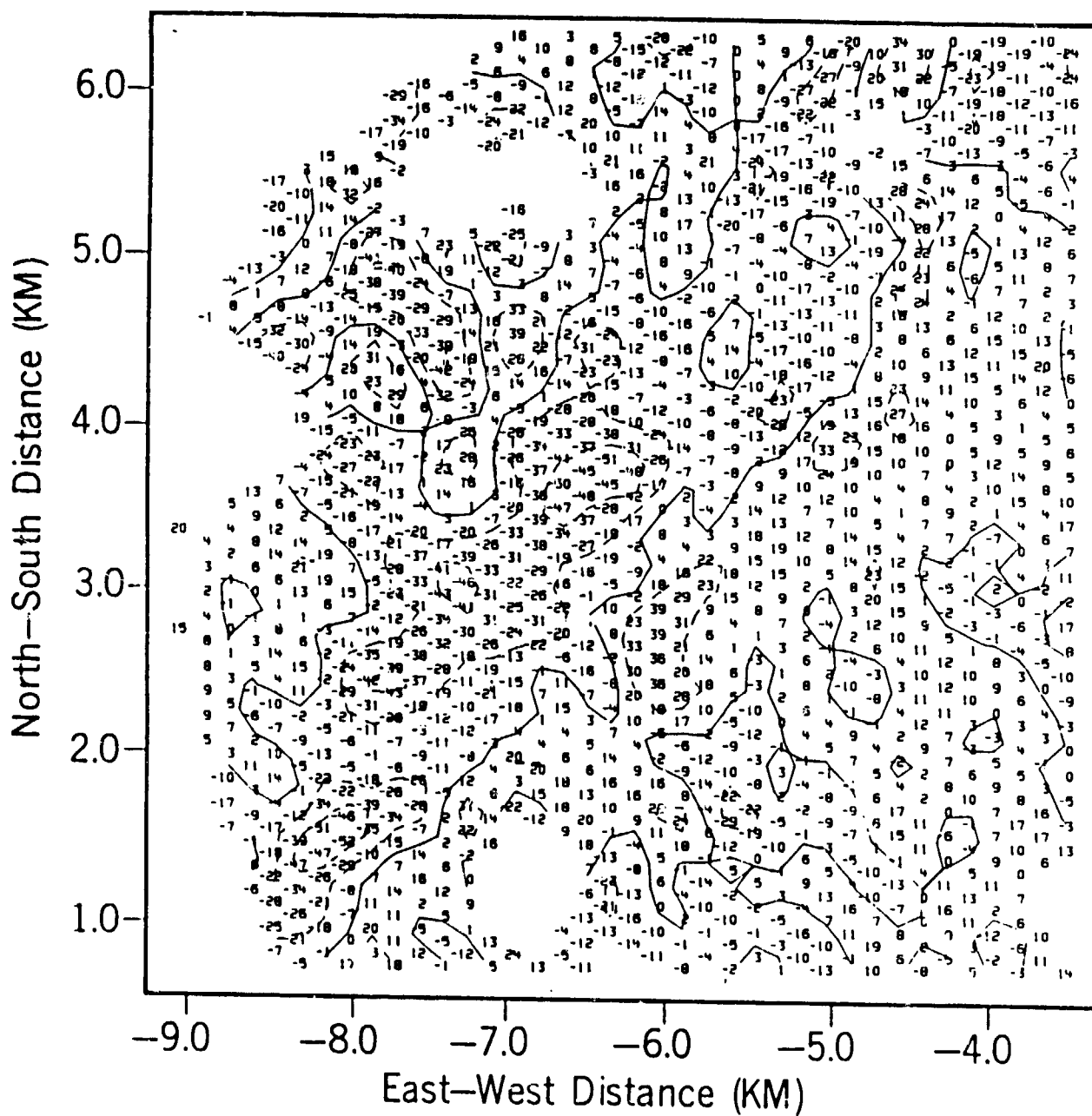


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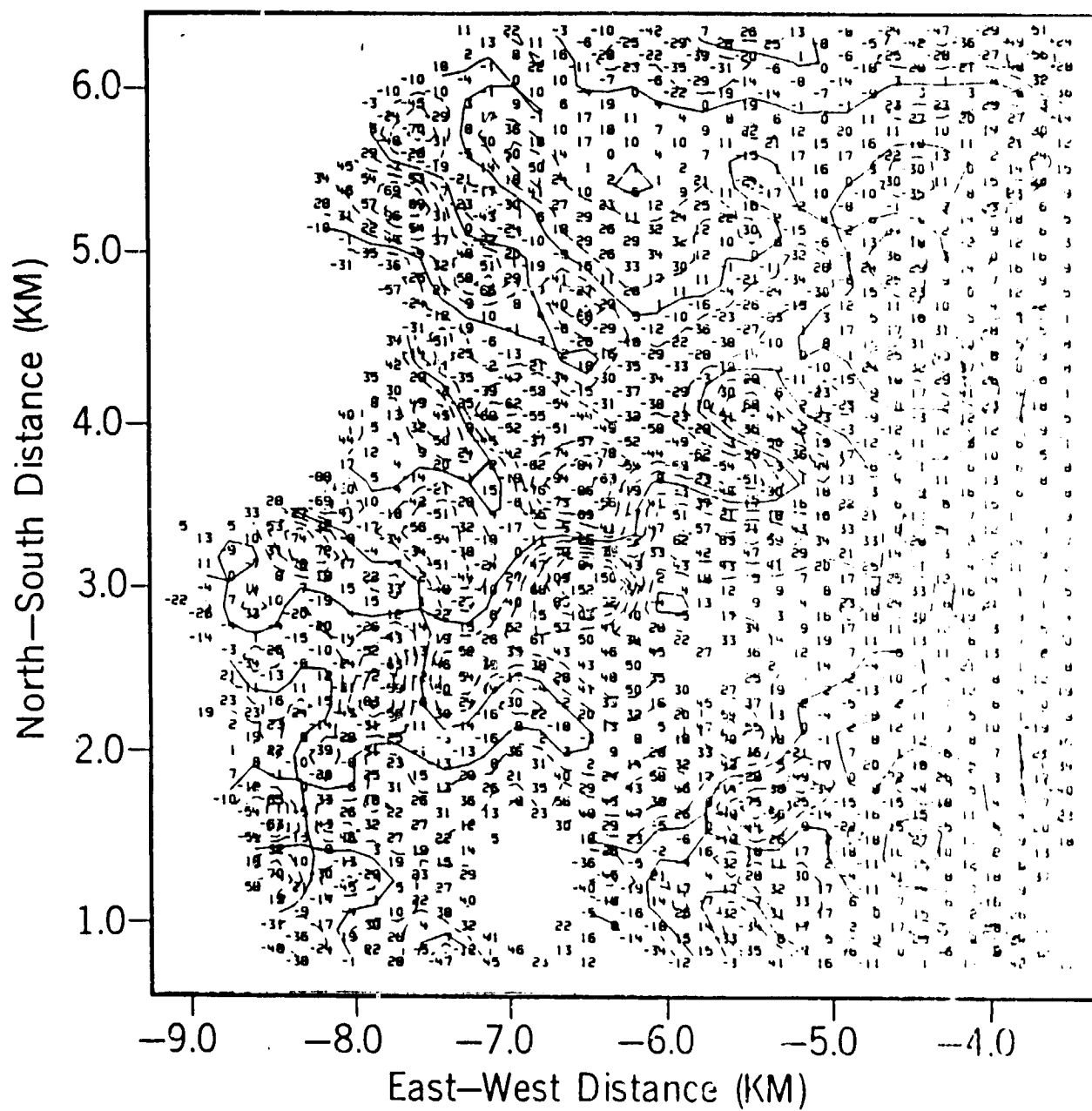
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Time 1254:04



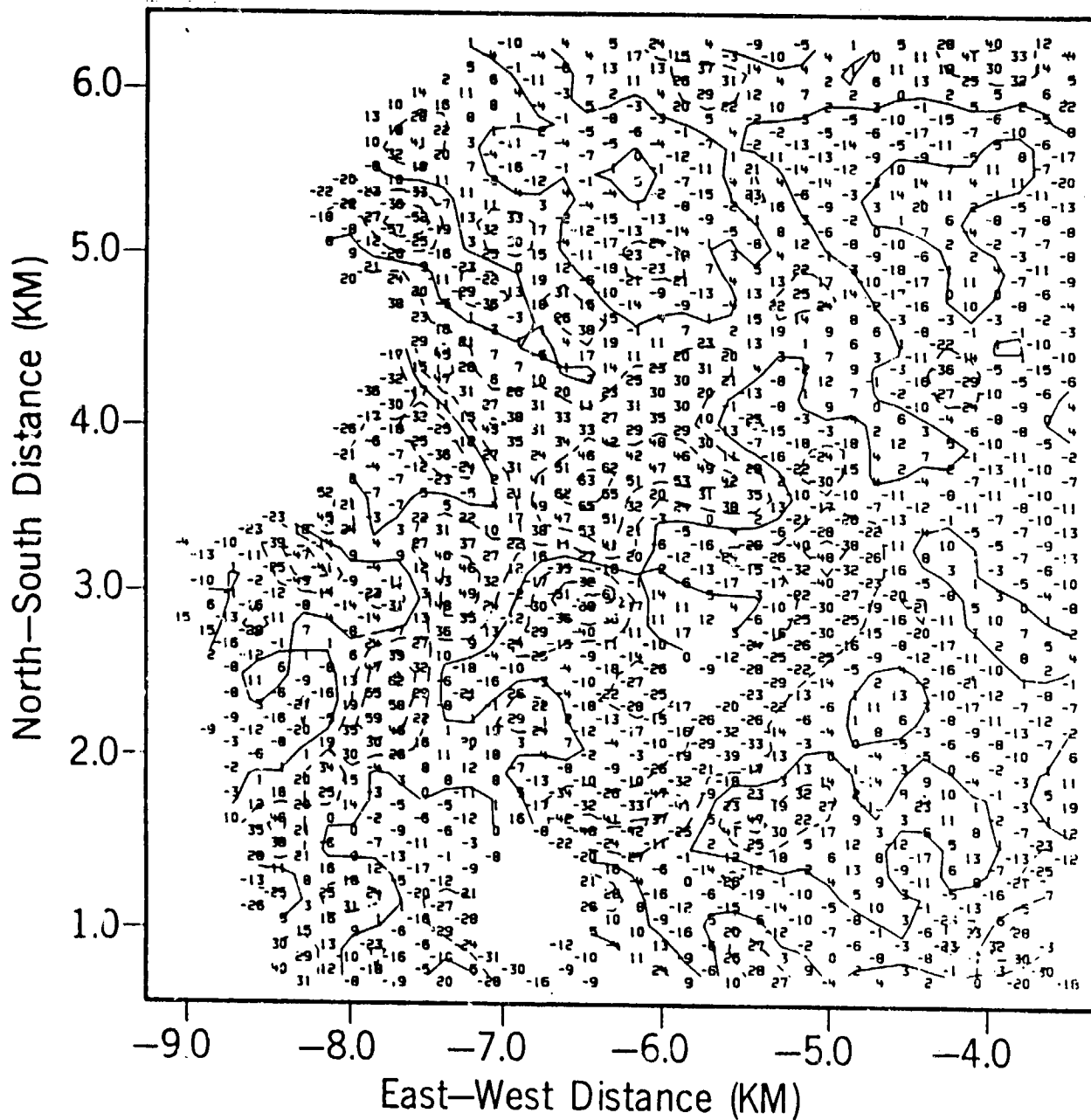
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Time 1255:44



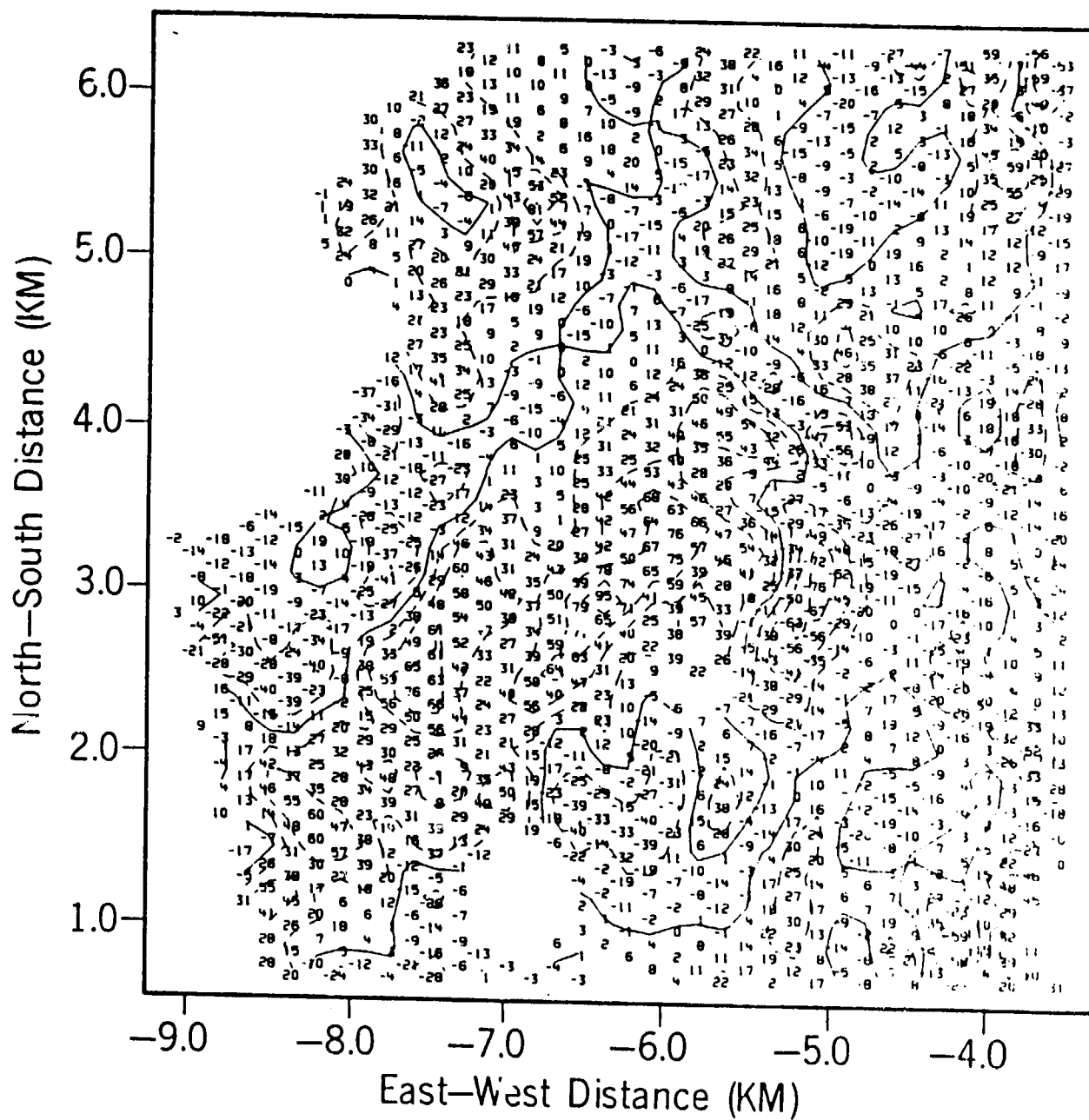
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Time 1255:44



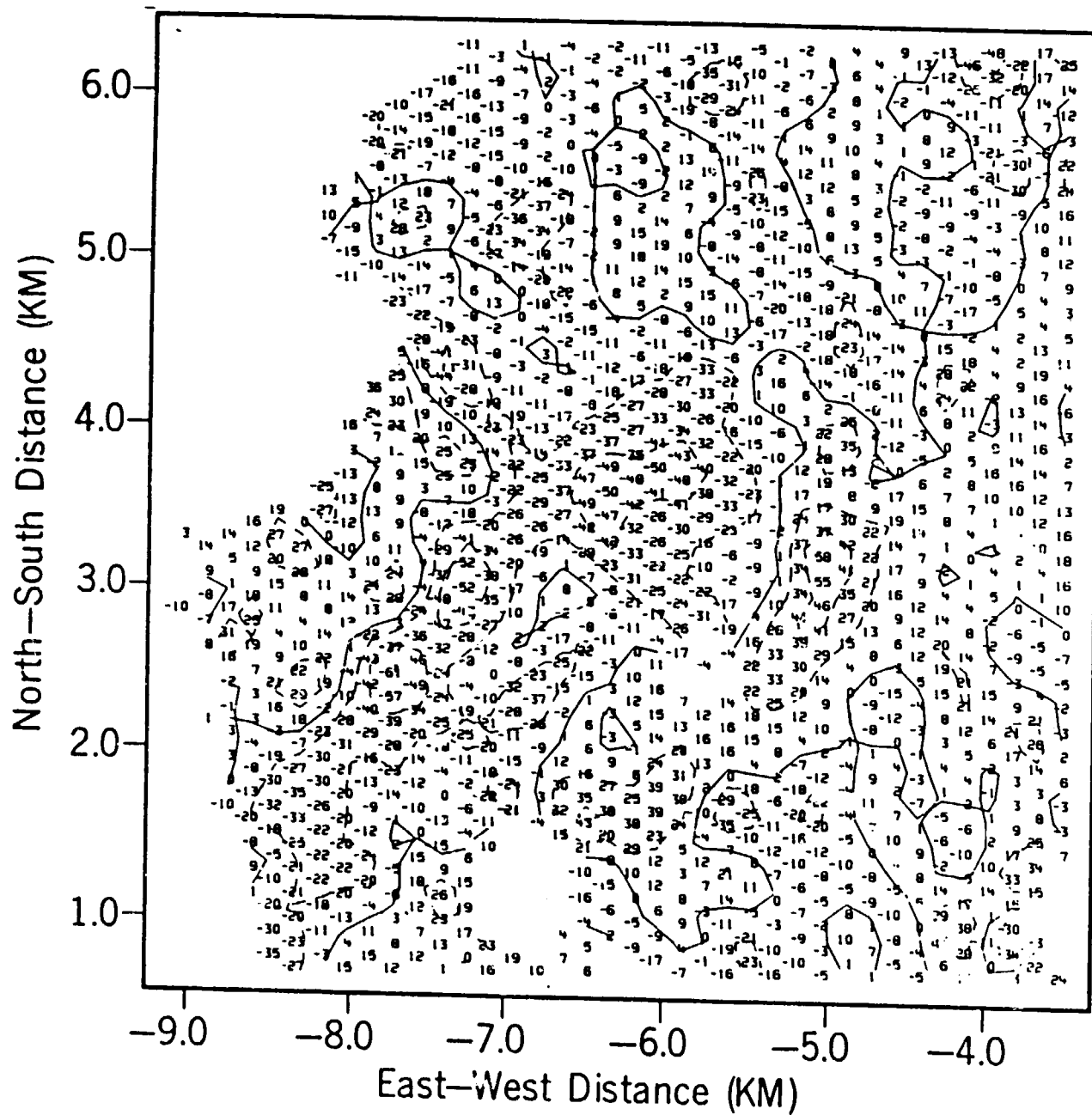
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Time 1255:44



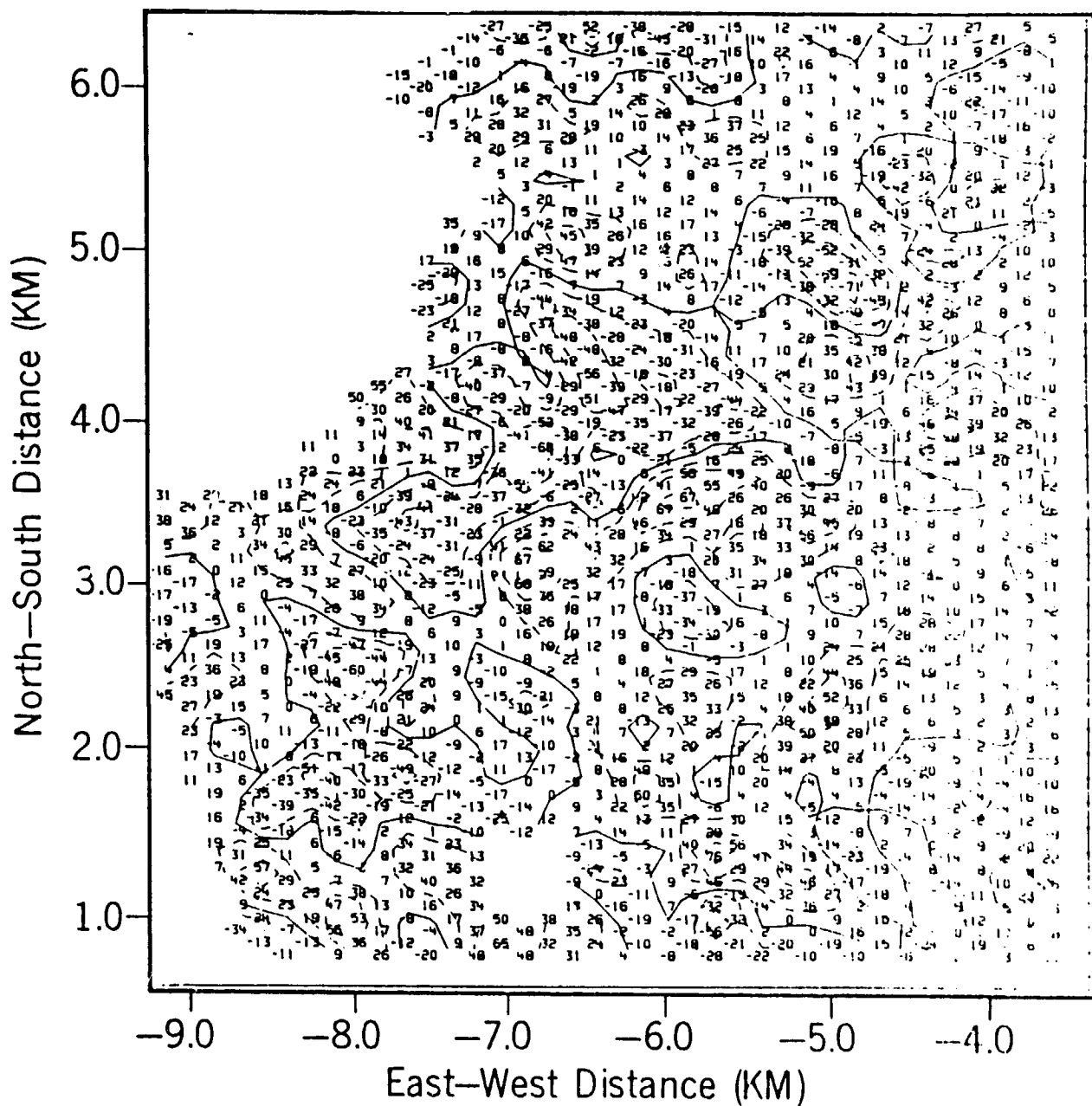
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Time 1255:44



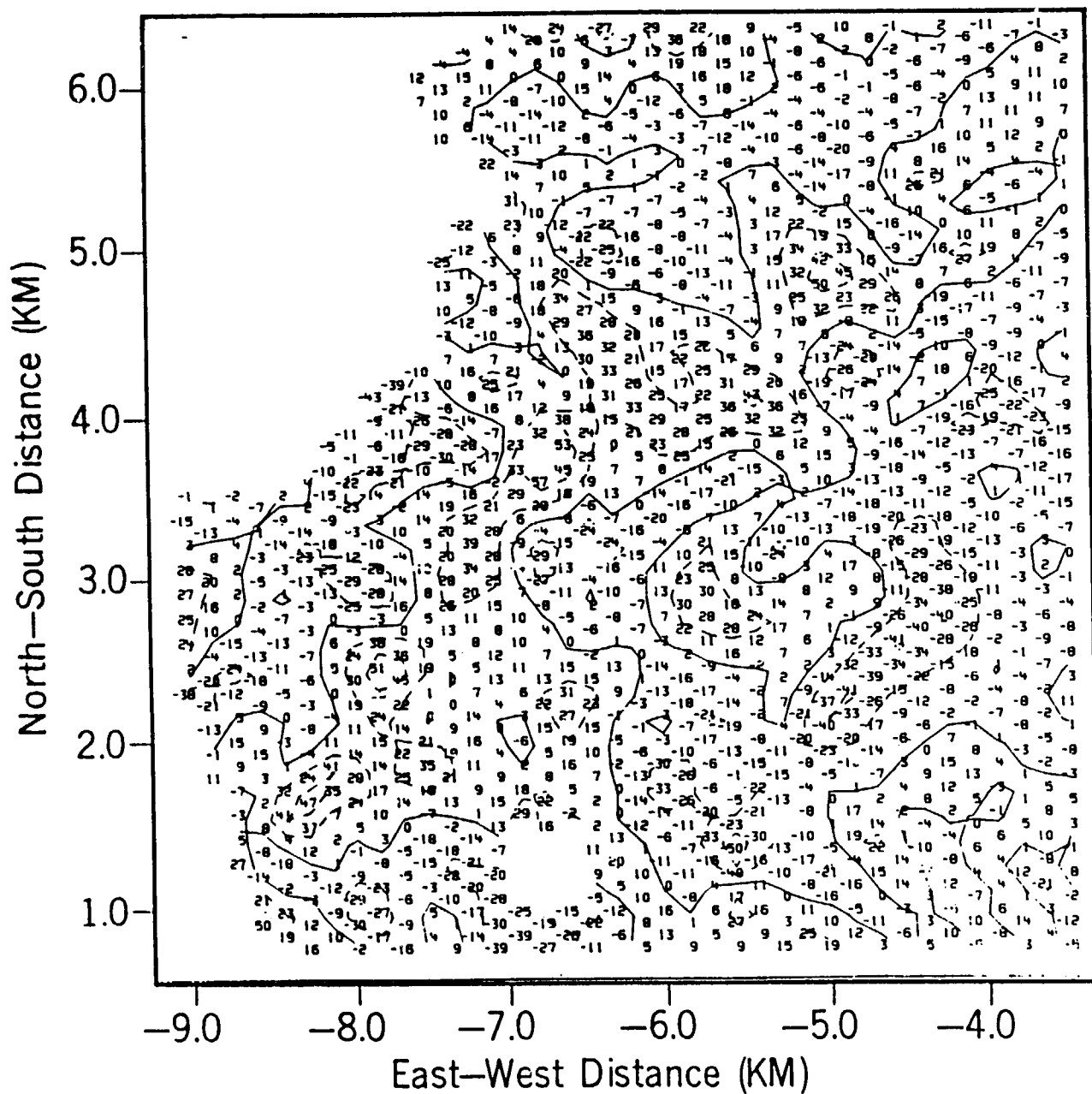
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Time 1257:24



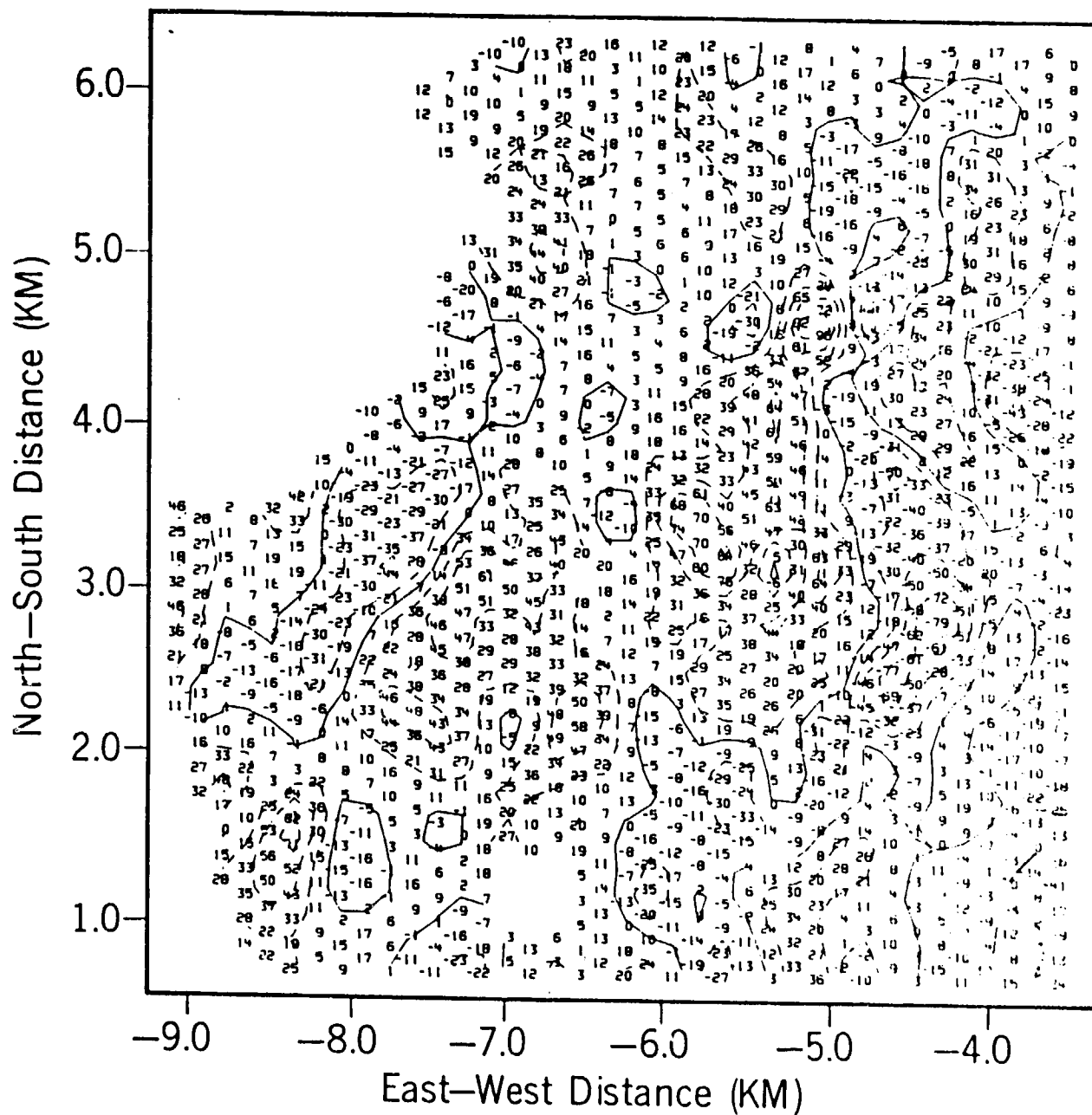
Field: FS X 1000

Time 1257:24



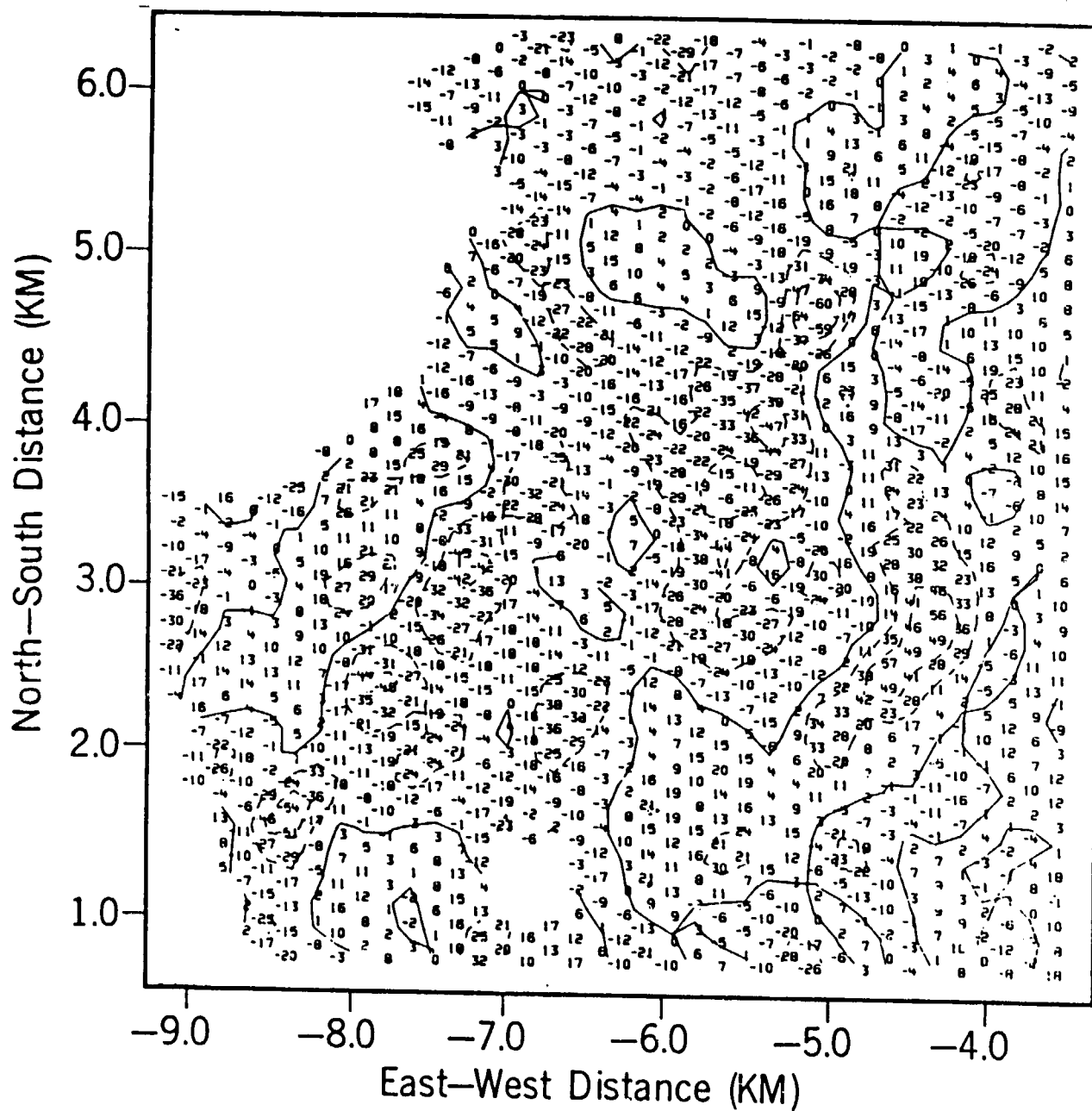
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Time 1257:24



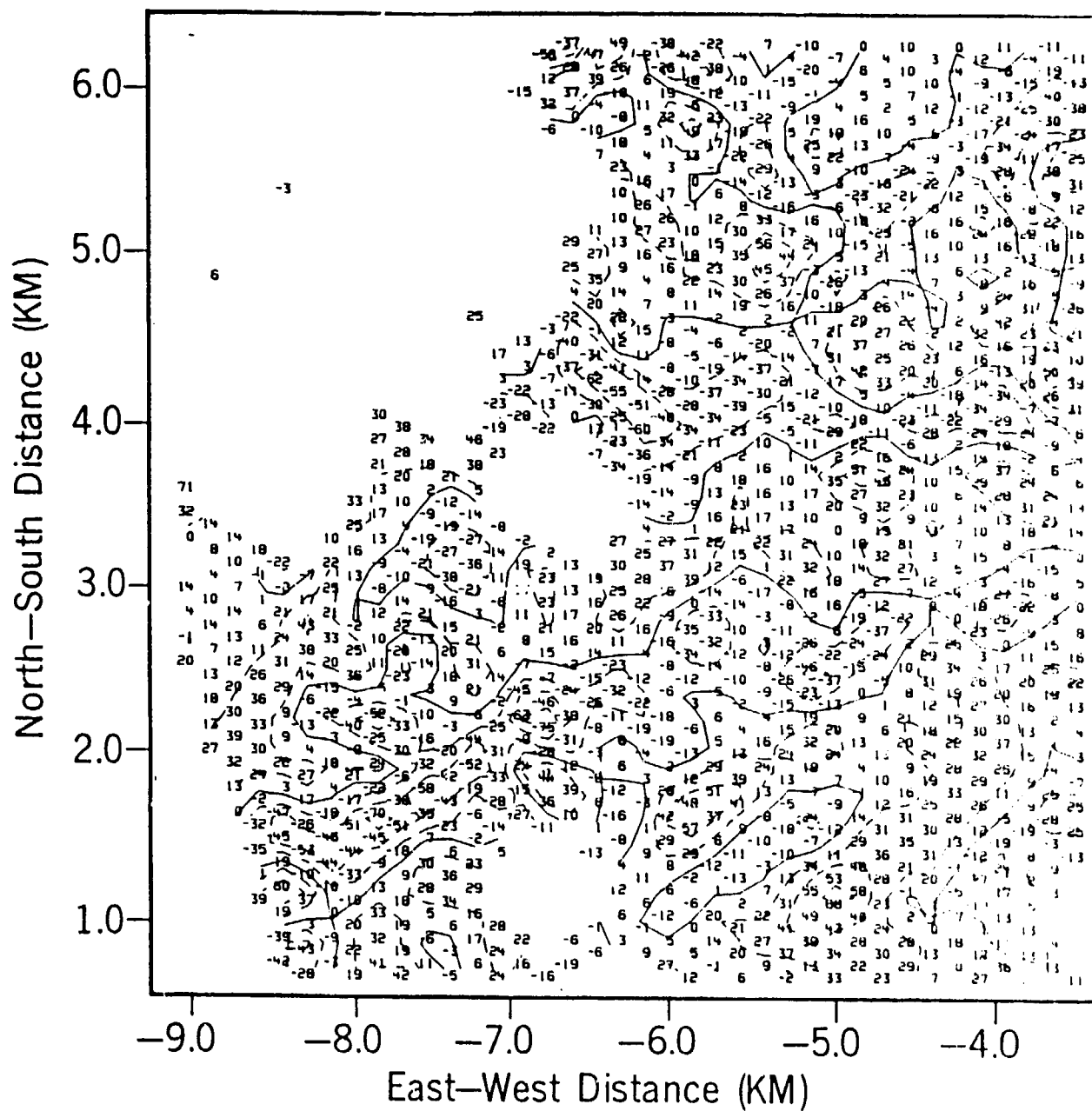
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Time 1257:24



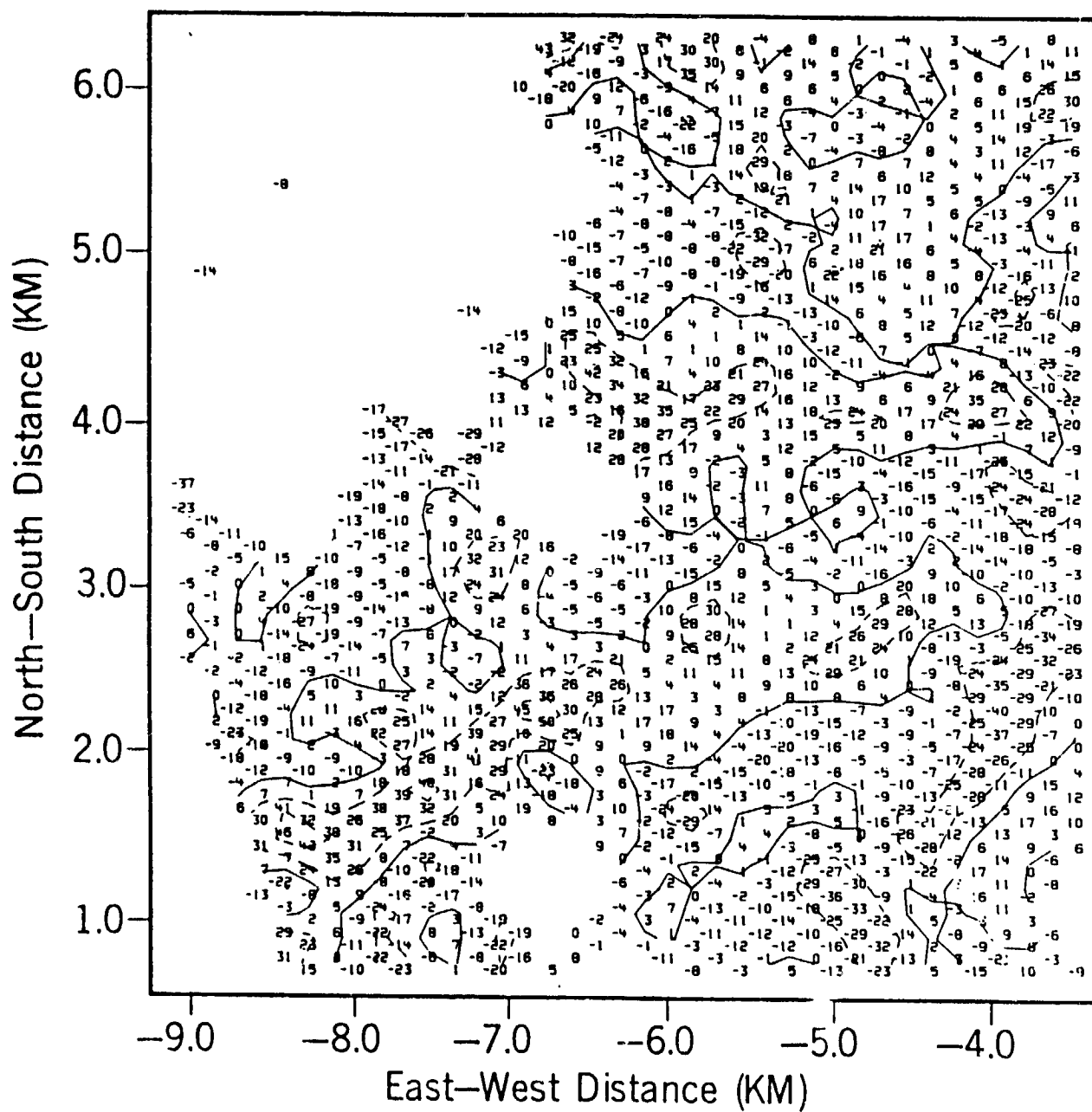
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Time 1259:04



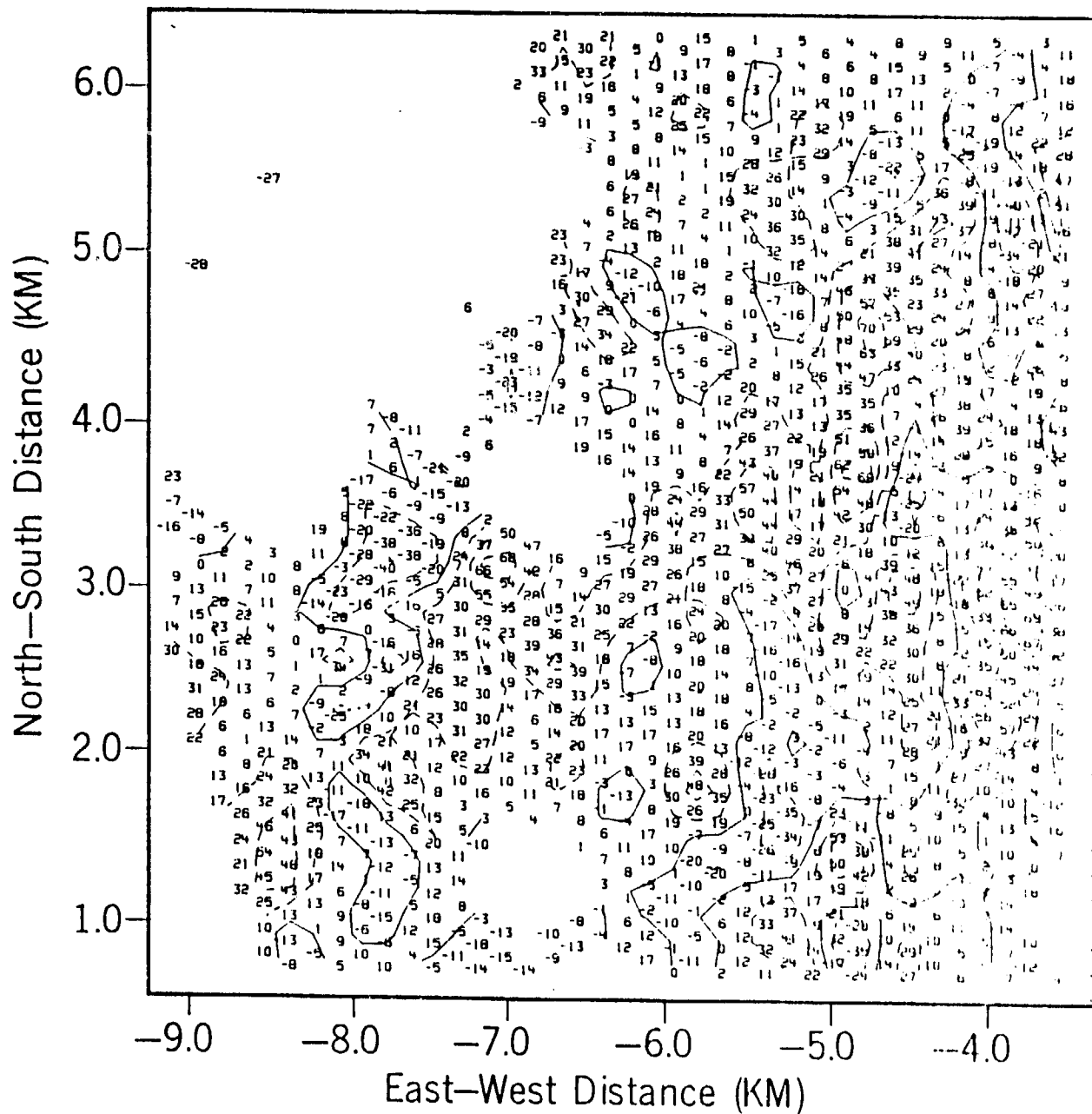
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Time 1259:04



Field: FE X 1000

Time 1259:04

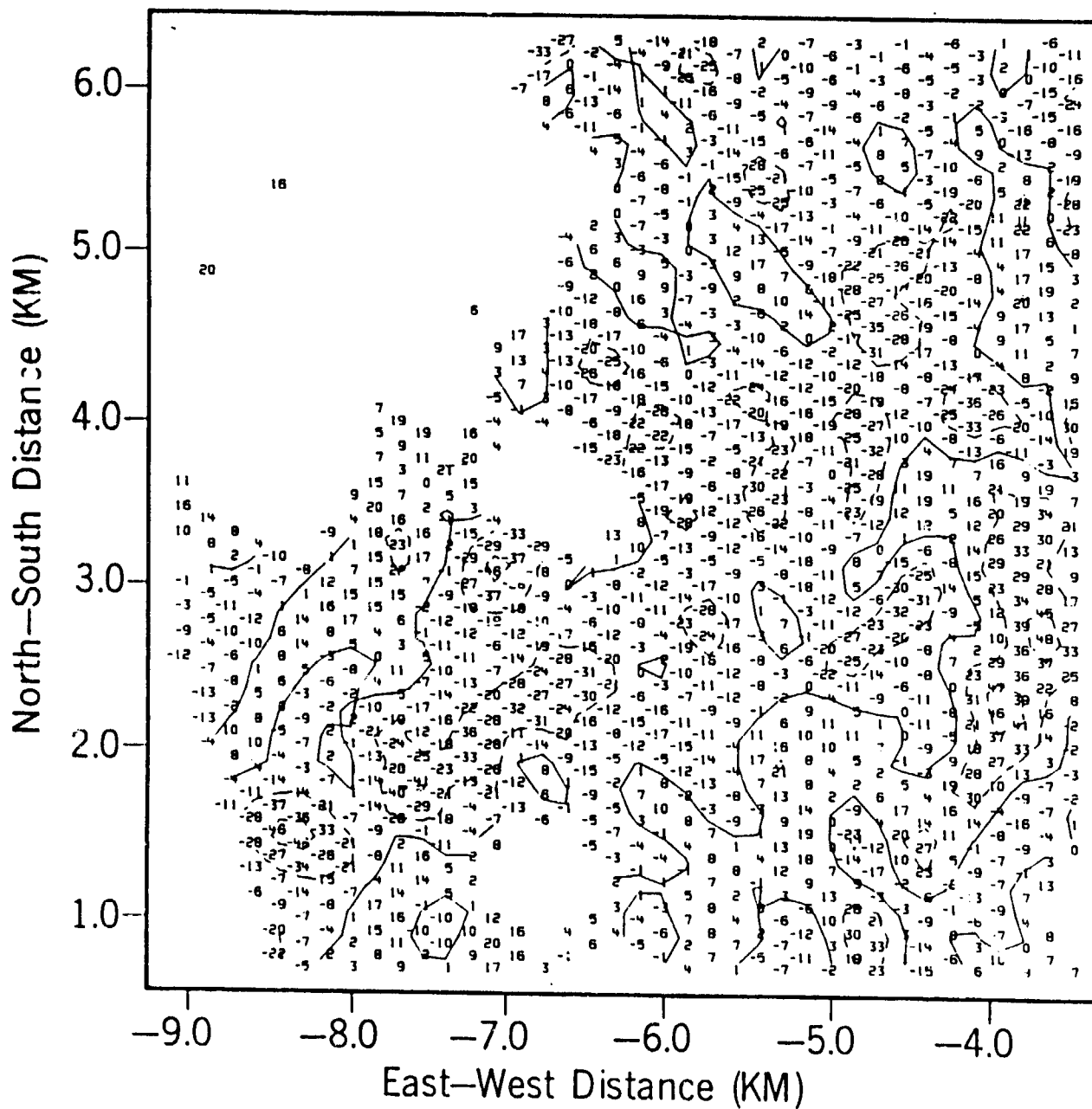


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Field: FW X 1000

Time 1259:04



APPENDIX III

Tape Format

TAPE HEADER

CARD IMAGE FORMAT (80 ASCII CHARACTERS/RECORD)
STRUCTURE

TAPE HEADER
[EOF]

VOLUME HEADER (VOLUME 1)

LINE 1 FORMAT(11X,A10,11X,A5,19X,F10.3)

TAPE NAME (A10) SCAN MODE (A5) MISSING DATA FLAG (F10.3)

LINE 2 FORMAT(13X,A10,15X,I3,15X,I3,I5)

VOLUME NAME (A10) VOLUME NUMBER (I3) PHYSICAL FILE (I3)(I5) FILES IN VOLM

LINE 3 FORMAT(19X,2I7,11X,2I7)

VOLUME BEGIN TIME (Y'MMDD HHMMSS) END TIME (2I7)

LINE 4 FORMAT(A)

TITLE FOR LINES 5 THRU 7

LINE 5-7 FORMAT(16X,A1,I6,3F10.3,A3,I7)

THESE LINES CONTAIN THE VOLUME COORDINATES (A1), COORDINATE INDICES (I6),
LIMITS, INCREMENT (3F10.3) AND DIMENSION (I7)
FOR EACH INDEX

LINE 8 FORMAT(15X,I2,28X,I8)

NUMBER FIELDS (I2) NUMBER POINTS/FIELD/PLANE (I8)

LINE 9 FORMAT(A)

FIELD NAMES IN SAME ORDER ON TAPE

LINE 10 FORMAT(1X,8A10)

10 CHARACTER FIELD NAMES

LINE 11 FORMAT(17X,I2)

NUMBERS OF RADARS (I2)

LINE 12 FORMAT(A)

RADAR COORDINATES

LINE 13 TO EOF FORMAT(7X,I2,3F8.3)

THESE LINES CONTAIN RADAR NUMBER AND COORDINATES(X,Y,Z) FOR EACH RADAR
[EOF]

EACH PLANE OF THE VOLUME IS WRITTEN IN A SEPARATE FILE

LINE 1 IS A FILE HEADER FORMAT(1X,2A10,F10.3,5I8)

CONTENTS OF FILE HEADER

TAPE NAME (A10)

VOLUME NAME (A10)

LEVEL OF PLANE (X,Y, OR Z) IN KM (F10.3)

NUMBER OF PLANES (FILES) IN VOLUME (I8)

PLANE (FILE) NUMBER THIS VOLUME (I8)

NUMBER OF FIELDS IN VOLUME (I8)

NUMBER OF POINTS/FIELD IN PLANE (I8)

PHYSICAL FILE NUMBER (I8)

THE FILE HEADER IS FOLLOWED BY THE FIRST FIELD FOR THIS PLANE. THE NUMBER OF
POINTS AND THE DIMENSIONS ARE GIVEN IN THE VOLUME HEADER.
THE DATA ARE WRITTEN WITH FORMAT(1X,6E13.6).

ALL FIELDS ARE WRITTEN CONSECUTIVELY IN THE ORDER GIVEN IN THE VOLUME HEADER
AND EACH FIELD BEGINS A NEW RECORD.

THE REST OF THE TAPE IS AS FOLLOWS

[EOF]

FILE HEADER (PLANE 2)

DATA ALL FIELDS

[EOF]

FILE HEADER (PLANE 3)

[EOF]

VOLUME HEADER (VOLUME 2)

TAPE NAME SCAN MODE CRT MISSING DATA FLAG .000
VOLUME NAME M311247 VOLUME NUMBER 1 PHYSICAL FILE 1 4 FILES IN VOLM
VOLUME BEGIN TIME 840531 124723 END TIME 840531 124725
VOLUME COORDINATE INDEX MIN MAX INCREMENT NUMBER
 X 1 -9.225 -3.375 .150 KM 40
 Y 2 .575 6.425 .150 KM 40
 Z 3 .075 .525 .150 KM 4
NUMBER FIELDS 4 NUMBER POINTS/FIELD/PLANE 1600
FIELD NAMES IN SAME ORDER ON TAPE
FN FS FE FW
NUMBER OF RADARS 2
RADAR COORDINATES
RADAR 1 .000 .000 .000
RADAR 2 .700 -6.070 .020

TAPE NAME SCAN MODE CRT MISSING DATA FLAG .000
VOLUME NAME M311249 VOLUME NUMBER 2 PHYSICAL FILE 6 4 FILES IN VOLM
VOLUME BEGIN TIME 840531 124903 END TIME 840531 124905
VOLUME COORDINATE INDEX MIN MAX INCREMENT NUMBER
 X 1 -9.225 -3.375 .150 KM 40
 Y 2 .575 6.425 .150 KM 40
 Z 3 .075 .525 .150 KM 4
NUMBER FIELDS 4 NUMBER POINTS/FIELD/PLANE 1600
FIELD NAMES IN SAME ORDER ON TAPE
FN FS FE FW
NUMBER OF RADARS 2
RADAR COORDINATES
RADAR 1 .000 .000 .000
RADAR 2 .790 -6.070 .020

TAPE NAME SCAN MODE CRT MISSING DATA FLAG .000
VOLUME NAME DREF1250 VOLUME NUMBER 3 PHYSICAL FILE 11 4 FILES IN VOLM
VOLUME BEGIN TIME 840531 125043 END TIME 840531 125043
VOLUME COORDINATE INDEX MIN MAX INCREMENT NUMBER
 X 1 -9.225 -3.375 .150 KM 40
 Y 2 .575 6.425 .150 KM 40
 Z 3 .075 .525 .150 KM 4
NUMBER FIELDS 4 NUMBER POINTS/FIELD/PLANE 1600
FIELD NAMES IN SAME ORDER ON TAPE
FN FS FE FW
NUMBER OF RADARS 2
RADAR COORDINATES
RADAR 1 .000 .000 .000
RADAR 2 5.130 6.530 -.060

TAPE NAME SCAN MODE CRT MISSING DATA FLAG .000
VOLUME NAME DREF1252 VOLUME NUMBER 4 PHYSICAL FILE 16 4 FILES IN VOLM
VOLUME BEGIN TIME 840531 125223 END TIME 840531 125223
VOLUME COORDINATE INDEX MIN MAX INCREMENT NUMBER
 X 1 -9.225 -3.375 .150 KM 40
 Y 2 .575 6.425 .150 KM 40
 Z 3 .075 .525 .150 KM 4
NUMBER FIELDS 4 NUMBER POINTS/FIELD/PLANE 1600
FIELD NAMES IN SAME ORDER ON TAPE
FN FS FE FW
NUMBER OF RADARS 2
RADAR COORDINATES
RADAR 1 .000 .000 .000
RADAR 2 5.130 6.520 -.060

